



Middle-to-Upper Palaeolithic site formation processes at the Bordes-Fitte rockshelter (Central France)

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ABSTRACT

Transformation in technological patterns associated with the Middle-to-Upper Palaeolithic transition between 50 and 40 ka in Western Europe and their relationship with the Neanderthal and Anatomically Modern Human populations and behaviors are issues that continue to stimulate heated debate. In this article we use the Middle and Early Upper Palaeolithic archaeo-stratigraphic record from the Bordes-Fitte rockshelter (les Roches d'Abilly site, Central France), a Bayesian analysis of the ages obtained by accelerator mass spectrometry radiocarbon on ultrafiltered collagen and by luminescence on quartz and feldspar grains, to establish a timeline for material culture and sedimentary dynamic changes during the Middle-to-Upper Palaeolithic transition. Technology, refitting studies and taphonomy of lithic artifacts recovered in the geoarchaeological field units D1 and D2 permit to characterize 3 reduction strategies (Levallois, Discoidal and Châtelperronian blade) that took place between the cold Heinrich events 5 and 4. We discuss the implications of the results to characterize the end of the Middle Palaeolithic, and for distinguishing anthropogenic and non-anthropogenic factors in Middle-to-Upper Palaeolithic assemblage's variability.

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1. Introduction

Changes in material culture associated with the Middle-to-Upper Palaeolithic transition between 50 and 40 ka in Western Europe, and their relationship with the Neanderthal and Anatomically Modern Human populations and behaviors are issues that continue to stimulate heated debate (Bordes, 1971; Howells, 1976; D'Errico et al., 1998; Mellars, 1996, 2005; Bar-Yosef and Bordes, 2010; Zilhão, 2013). Two main characteristics of the geoarchaeological record of this period could explain the difficulty in elaborating a reliable timeline with a resolution sufficient to understand material culture changes observed during this period.

Firstly, the key Middle-to-Upper Palaeolithic (MUP) sequences of Western-south France were affected by successive cold Heinrich events (HE) (Heinrich, 1988) that caused a periglacial environment

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along the Atlantic continental margin (Sanchez-Goni and Harrison, 2010), propitious to erosive processes attested by stratigraphical discontinuities and post-depositional disturbances in archaeo-stratigraphic sequences (Bertran et al., 2010; Goldberg and Macphail, 2006). Lithic assemblage technology and typology (Rigaud, 1996), refitting between lithic (Bordes, 2003; Bachellerie, 2011) or faunal remains (Morin et al., 2005; Discamps et al., 2012), surface damage analysis correlated with technological study of lithic remains (Zilhão et al., 2006; Villa and Soressi, 2000; Soressi, 2011), and study of patterns in the orientation of elongated artefacts (Bertran et al., 1997; Lenoble and Bertran, 2004; McPherron, 2005) have been applied to evaluate mixing disturbances and stratigraphic coherence for MUP archaeological assemblages. However, these taphonomic analyses have been realized independently of the reconstruction of sedimentary and pedogenic processes, and have yet to be applied systematically to the key MUP sequences in order to establish the relationship between dated samples and occupation by human or carnivores (Zilhão, 2013).

Secondly, the limitations of ^{14}C dating beyond 35 ka cal BP have been proven by studies which have re-dated sites using AMS radiocarbon dating with ultrafiltration. These have systematically revealed older results compared to those obtained when other pretreatment chemical methods were applied (Higham et al., 2006; Jacobi et al., 2006). Despite some controversies (Hüls et al., 2007; Hublin et al., 2012), using this sample preparation protocol is leading to the creation of a new chronological framework for the Late Middle Palaeolithic (LMP) and Early Upper Palaeolithic (EUP) in Western Europe (Gravina et al., 2005; Hublin et al., 2012;

Higham, 2011; Higham et al., 2010, 2011, 2014; McPherron et al., 2012; Richter et al., 2013; Talamo et al., 2012). Until very recently, there are still few cases of application of the ultrafiltration protocol to date occupations corresponding to the very end of the Middle Palaeolithic in Europe (Higham et al., 2014), and most of the contexts attributed to this period were based on lithic technology, faunal biozonation or their relative position in archaeological stratigraphic sequences (Jaubert et al., 2011).

Les Roches d'Abilly is an archaeological site located in Central France, along the Creuse Valley, that consists of small, horizontal

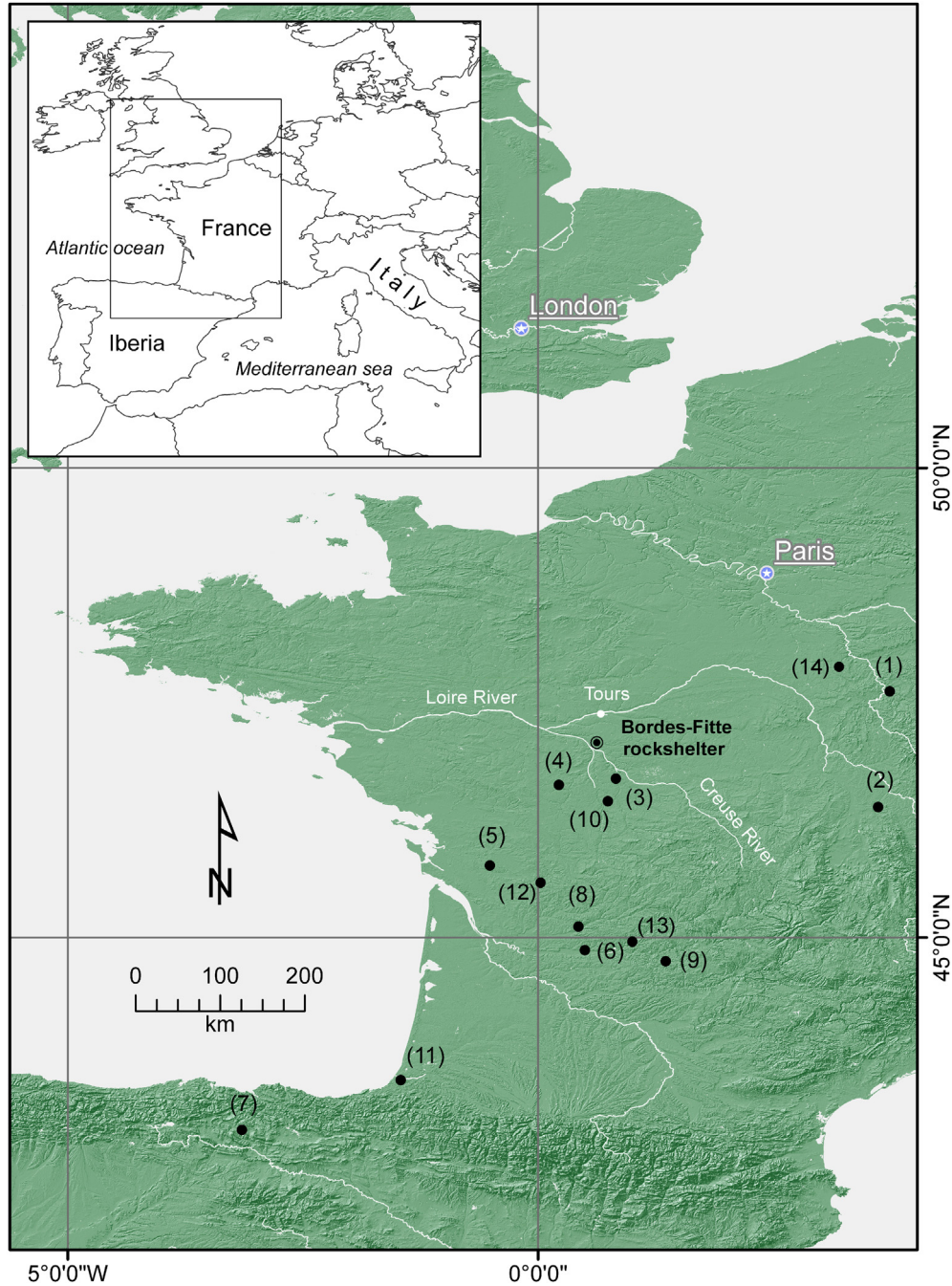


Fig. 1. Location map of the Châtelperronian and Late Middle Palaeolithic sites mentioned in the text, and Bordes-Fitte rockshelter (LRA-BF). (1) La Grotte du Renne, Grotte du Bison and Galerie Schoepflin (Arcy-sur-Cure), (2) Grotte des Fées (Châtelperron), (3) Grotte des Cottés (Saint-Pierre-de-Maillé), (4) Grotte de Grande Roche (Quincy), (5) La Roche à Pierrot (Saint-Césaire), (6) Canaule II (Creyse), (7) Labeko Koba (Arrasate), (8) La Côte (Neuvic), (9) Roc-de-Combe (Payrignac), (10) Les Rochers-de-Villeneuve (Lussac-les-Châteaux), (11) Le Basté (Saint Pierre d'Irube), (12) La Quina (Gardes-Pontaroux), (13) Le Moustier (Peyzac-le-Moustier), (14) Les Bossats (Ormesson).

inter-bedded rockshelters and caves, which are exposed to the south and southwest along a quarried escarpment at the right margin of the valley, at the southern limit of the Touraine region (Aubry et al., 2012). It is midway between La Roche à Pierrot (Saint-Césaire) and the Grotte du Renne (Arcy), which are two key MUP occupational sequences that continue to stimulate heated debate (Fig. 1, Higham et al., 2010, 2014; Hublin et al., 2012).

Since 2007, excavations have been carried out at several locations along the ~300 m long cliff where large collapsed roofs and walls slabs of compact calcarenite and/or silicified have preserved siliciclastic deposits containing Middle Palaeolithic (MP) and EUP remains (Aubry et al., 2012, Fig. 2). The analysis of the difference in bleaching rates of quartz and feldspar luminescence signals indicates that quartz and feldspar grains of these deposits are well-bleached and ages obtained by luminescence are coherent with ^{14}C measurements (Buylaert et al., 2012; Murray et al., 2012).

The data presented in this article concern only the Bordes-Fitte rockshelter (LRA-BF), taking into account archaeological data concerning level C, rockshelter collapse events and stratigraphical subdivisions of the level D established in recent excavations (years

2012–2013). The results include systematic refitting of lithic remains, micromorphological study, new AMS radiocarbon ages obtained on ultrafiltered bone and tooth gelatine, and quartz data from Aubry et al. (2012) which were combined with additional measurements (Thomsen et al., submitted for publication). These data have allowed for the creation of a refined timeline for anthropogenic and non-anthropogenic processes that occurred at the time of the Neanderthal to Anatomically Modern Human transition.

2. Material and methods

Considering the two main problems defined above (integrity of geoarchaeological context and dating methods) to explain the difficulty in defining a reliable timeline for the changes through the MUP transition, our approach relies on the reconstruction of the Bordes-Fitte rockshelter site formation processes and to establish a more accurate timescale, correlating AMS ^{14}C and luminescence results.

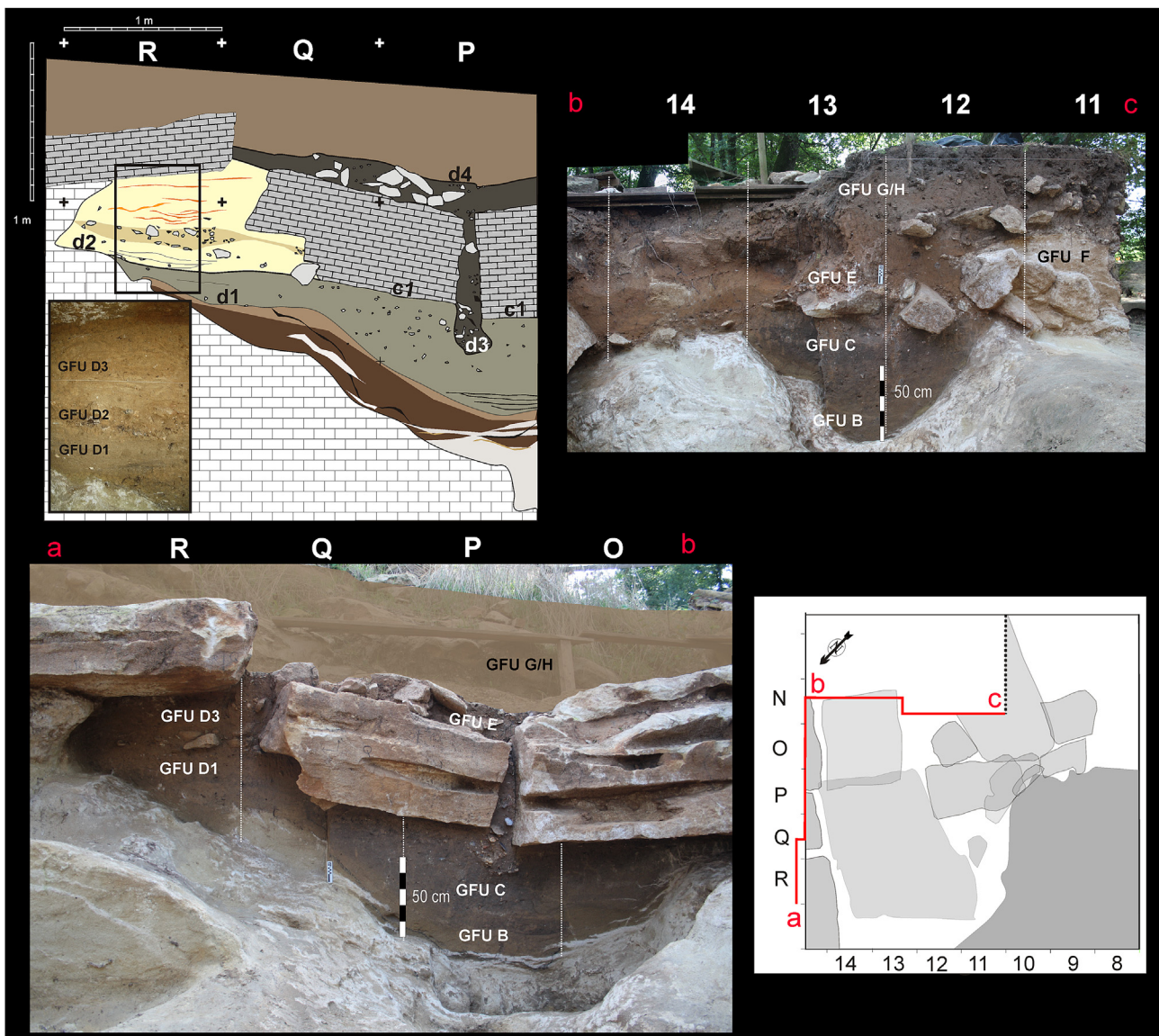


Fig. 2. Perpendicular sections in the LRA-Bordes-Fitte collapsed rockshelter deposits with indication of non-depositional unconformity and disconformity (d#), rockshelter roof and wall collapse event (c#), and Geological Field Unit B to G/H.

2.1. Geoarchaeological approach

The data were collected under a standard geoarchaeological fieldwork approach including: geomorphological and geological study of the site surroundings, description and stratigraphic correlation of the rockshelter siliciclastic infill, sampling for micromorphological analysis and dating. Field work included the systematic description of cross-sections and profiles to reconstruct stratigraphic events, their vertical and lateral variations and archaeological remains analysis. Geoarchaeological field units (GFU) were identified on the basis of lithostratigraphic, pedological and archaeological criteria and used as field categories. The field units were later grouped into geoarchaeological complexes (GC), taking into consideration only major changes in the depositional style, as well as major unconformities that would imply erosional phases (disconformities) or long hiatuses (details in [Angelucci, 2002](#)).

The samples for micromorphological observation were collected at the end of the 2008 and 2012 field seasons. The thin section prepared by Thomas Beckmann Laboratory (Germany) and the EPOC-UMR 5805 Laboratory (France), were observed under a polarizing Stereo-microscope (Nikon) with magnifications between 20× and 140× using plane-polarized light (PPL) and cross-polarized light (XPL). Description of the thin sections follows the guidelines proposed by [Bullock et al. \(1985\)](#), [Stoops et al. \(2010\)](#) and [Stoops \(2003\)](#), with additions from [Courty et al. \(1989\)](#) and [Goldberg and Macphail \(2006\)](#).

The field and laboratory sediment's descriptions were made using a comprehensive form addressing the sedimentary, pedogenic and anthropogenic characteristics of the deposits ([FAO-Isric, 1990](#); [Goldberg and Macphail, 2006](#); [Keeley and Macphail, 1981](#)) to collect the data necessary to assess the site's formation processes.

2.2. Analysis of lithic industries

The lithic assemblages were systematically studied for weathering and displacement damage, documenting the breaking, edge and surface abrasion, and the degree of patination. The origin of the raw material was determined by a macroscopical comparison with hand samples collected in the different geological formations of the Creuse and Cher River basins ([Aubry, 1991](#)), and the characterization of non-local flint was complemented by a binocular examination of the micro-texture and of the fossil content. The lithic remains were studied using a typo-technological approach, in order to assess reduction sequences and to define retouched tool types ([Boëda, 1993, 1994](#); [Dibble and Bar-Yosef, 1995](#); [Pelegrin, 1995](#)), distinguish coherent assemblages, and reconstruct sedimentary dynamic and post-depositional processes considering spatial analysis, systematic refitting ([Hofman and Enloe, 1992](#); [Villa, 1982](#)), stratigraphical relationship and three-dimensional plotting using ArcScene 10.

2.3. Radiometric dating

The chronological data were obtained by Accelerator Mass Spectrometry (AMS) ^{14}C and Optically Stimulated Luminescence (OSL) multi-grain dating, the latter using quartz and feldspar. Several small fragments of wood charcoal were found in the GFU C, but dating at the Oxford Laboratory failed due to low yield. Thus, all of the samples submitted for radiocarbon dating were of compact mammal bone or tooth identified to species level by Laure Fontana. The Beta Analytic protocol of bone pretreatment is collagen extraction with alkali applying the Longin method ([Longin, 1971](#)). The bone samples dated at the Lyon and Oxford Radiocarbon

Accelerator Unit (ORAU) were based on chemical pretreatment, target preparation and AMS measurement procedures outlined in previous publications ([Higham et al., 2006](#); [Brock et al., 2010](#)).

For the luminescence multi-grain dating, purified quartz and K-rich feldspar grains (180–250 μm) were extracted from the sediment in the usual manner ([Aitken, 1985, 1998](#)). The dose in quartz was measured using a double SAR procedure ([Murray and Wintle, 2000](#); [Banerjee et al., 2001](#)) employing infrared (IR, 870 \pm 40 nm) stimulation prior to blue (470 \pm 30 nm) stimulation (both at 125 °C for 40 s) and detection in the UV (Hoya U-340). Thermal treatment used a preheat of 260 °C for 10 s, a cutheat of 220 °C and an elevated temperature (280 °C) 40 s blue light stimulation after every SAR cycle ([Murray and Wintle, 2003](#)). Individual aliquots (~5 mm diameter) consisted of a monolayers of grains mounted on stainless steel cups. The usual laboratory tests were undertaken to confirm the appropriateness of the protocol (e.g. recycling, IR depletion ratio, recuperation, dose recovery). The initial 0.5 s of the signal minus an early background (interval 0.5 s–1.0 s) was used for all quartz calculations. The dose in feldspar was measured using IR stimulation and detection in a blue-violet window (Schott BG39/Corning7-59 filters) with 2 mm diameter grain monolayers in stainless steel cups. A post-IR IR SAR protocol ([Thomsen et al., 2008](#); [Buylaert et al., 2012](#); [Thiel et al., 2011](#)) was used with a 60 s 320 °C preheat, first IR stimulation at 50 °C (200 s) and the second stimulation (200 s) at 290 °C and a high-temperature (325 °C) IR exposure (200 s) after each SAR cycle. [Thiel et al. \(2011\)](#), [Thomsen et al. \(2011\)](#) and [Buylaert et al. \(2012\)](#) have shown that this post-IR IR signal is stable – i.e. anomalous fading ([Spooner, 1994](#)) is not detectable. Calculations were made using the initial 2 s of the post-IR IRSL decay curve minus a background derived from the last 50 s.

Dose rates were determined using high-resolution gamma spectrometry ([Murray et al., 1987](#)) and the conversion factors of [Guérin et al. \(2011\)](#). For K-feldspar extracts the internal beta dose rate of ^{40}K was calculated assuming an effective K concentration of 12.5 \pm 0.5% ([Huntley and Baril, 1997](#)). Additional OSL single-grain measurements for four quartz samples, with doses in the range ~60–110 Gy, have revealed that multi-grain analysis gives the most accurate ages (for details see [Thomsen et al., submitted for publication](#)). The radiocarbon determinations obtained by the ultrafiltration protocol and luminescence ages were used to construct a model for the Bordes-Fitte rockshelter site using Bayesian software OxCal 4.2 ([Bronk Ramsey, 2013](#)), and the INTCAL13 curve ([Reimer et al., 2013](#)).

3. Sedimentary sources and dynamics in the Bordes-Fitte rockshelter

The exposed bedrock along the Roches d'Abilly cliff face is composed of an Upper Cretaceous (Upper Turonian) lithostratigraphic succession of chalk, calcareous tufa with some nodular chert, sandy marl, and alternating clastic limestone (bio-calcirudite and calcarenite affected by decalcification and silicification) with thin chert bands, topped by compact silicified limestone ([Aubry et al., 2012](#); [Médioni et al., 1974](#)). The sedimentary cover includes Eocene cherty mudstone and lacustrine marls and limestone, Mio-Pliocene formations of mudstone units that overlie the Eocene deposits in some locations, and Quaternary alluvial deposits with some quartz and flint pebbles.

During the final episode of the Upper Cretaceous sub-aerial exposure, and throughout Quaternary times, the carbonate bedrock was affected by several weathering processes including dissolution ([Laignel et al., 1999](#)), silicification, ferruginization, and freeze-thaw. The dissolution and freeze-thaw mechanisms were responsible for the development of an important secondary

porosity that resulted in the development of karstic caves and rockshelters, which were subsequently filled by siliciclastic and carbonate sediments (Aubry et al., 2012).

From the geoarchaeological field description of the sediments in the LRA-BF rockshelter (Aubry et al., 2012, Figs. 2 and 3), it is clear that the siliciclastic inputs are related to the lithology of the incised Cenozoic overlying units (through their sub-aerial erosion and alluvial or down-slope transport and redeposition) coupled with weathering of the carbonate bedrock. As indicated by the common presence of quartz grains, the ultimate origin of the siliciclastic fraction is the widespread Mio-Pleistocene detrital units found lying on the Upper Cretaceous bedrock. The clay and silt inputs, however, are probably also derived from the erosion of earlier Eocene units and of pre-existing cave/rockshelter sediments. Chemical weathering cannot be excluded in the formation of part of this fine fraction.

Carbonate rock fragments of various sizes are present throughout the geoarchaeological sequence (Fig. 3). The origin of these rocky carbonate fragments most likely resides in the normal processes of mechanical degradation of the roof and walls of the caves/rockshelters, as indicated by their heterometry and varied shape. However, it cannot be excluded that some of the smaller-sized carbonate rock fragments (clasts between 2 and 4 cm in size) were brought into the karstic forms from the outside, or are

related to freeze-thaw weathering of roofs and walls near the entrance, followed by rapid short-distance displacement due to water flow. Such hypothesis could be confirmed by the clast's sharp edges morphologies observed during the micromorphological analysis of the GFU D2. Layered coatings of unsorted silts and many silty clay cappings on the upper part of clasts are other indicators of possible freeze-thaw impact, observed in the GFU D2 (Fig. 5 and Table 1). These features, particularly the cappings, are traditionally related with melt water percolation during seasonal surface soil thaw (Van Vliet Lanoë, in Stoops et al., 2010). The water logging resulting of a seasonal soil thaw or another cause not identified, could also explain the partial ground mass collapse of the GFU D1 micro-facies (Fig. 5).

In terms of sedimentary dynamics, the lithostratigraphic sequences are composed of near-surface sedimentary facies with vertical and lateral variations within a context dominated by run-off and gravitational processes. There is evidence for several episodes of sediment slope-wash, alluvial and lacustrine deposits, as well as hiatuses in sediment deposition and main erosional phases (Figs. 3 and 4 and Table 1).

After/during the opening of an inter-bedded cave (or already a rockshelter with the entrance at the local hydrographic base level), the carbonate bedrock started to weather (GFU A) where it was not covered by sediments, and perhaps this process continues

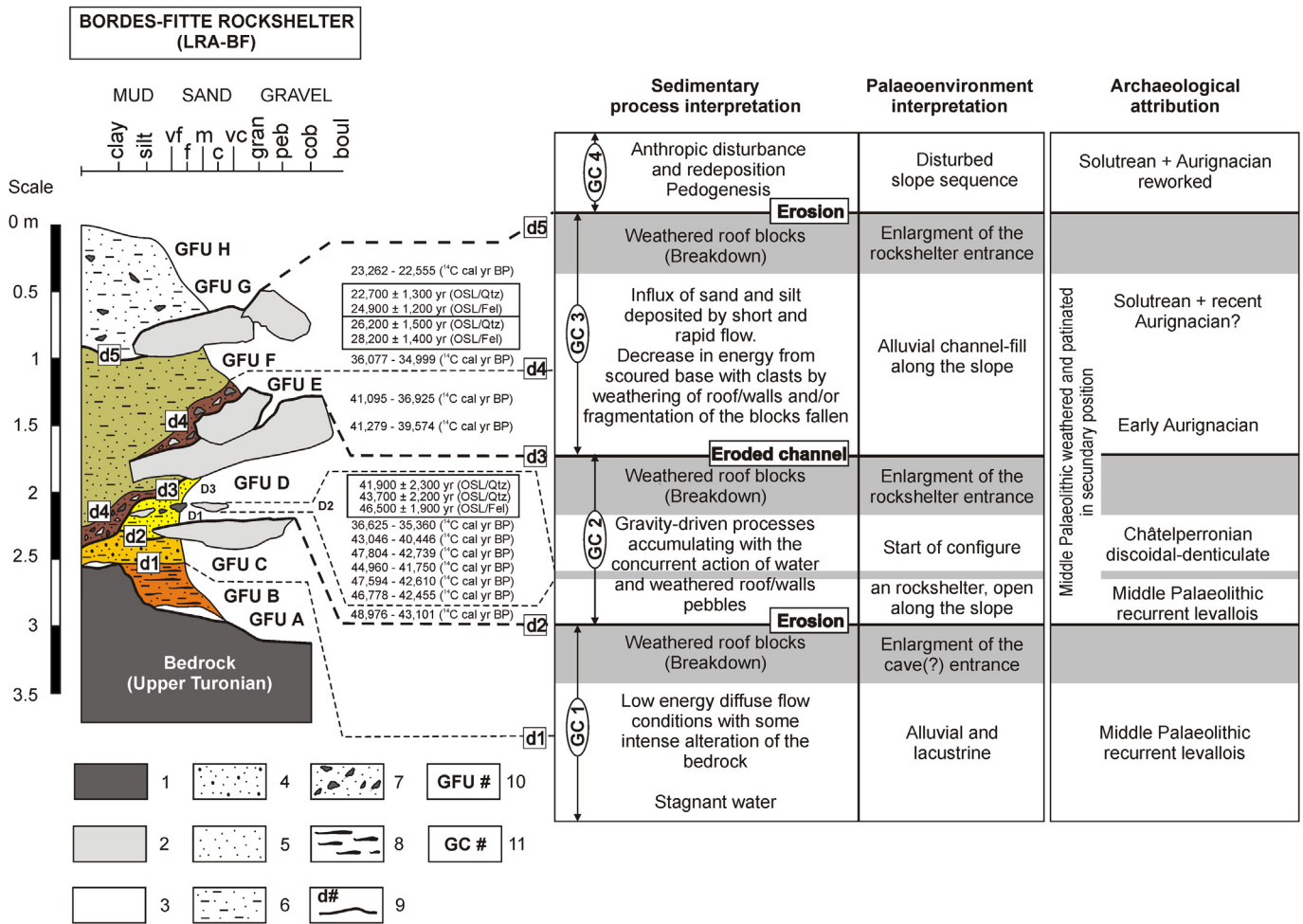


Fig. 3. Les Roches d'Abilly Bordes-Fitte rockshelter synthetic log showing the filled stratigraphic sequence, sedimentary process interpretation, palaeoenvironmental interpretation and archaeological attribution (modified from Aubry et al., 2012). Bio-calciuridite and calcarenite (1), silicified limestone (by roof rockshelter) (2), bioturbated calcareous sand (3), polymodal siliciclastic sand (4), unimodal siliciclastic sand (5), siliciclastic sandy-silt-clay (6), matrix-supported breccia (7), organic matter layers (8), non-depositional unconformity and disconformity (erosional unconformity) (9), geoarchaeological field unit (10), geoarchaeological complex (11).

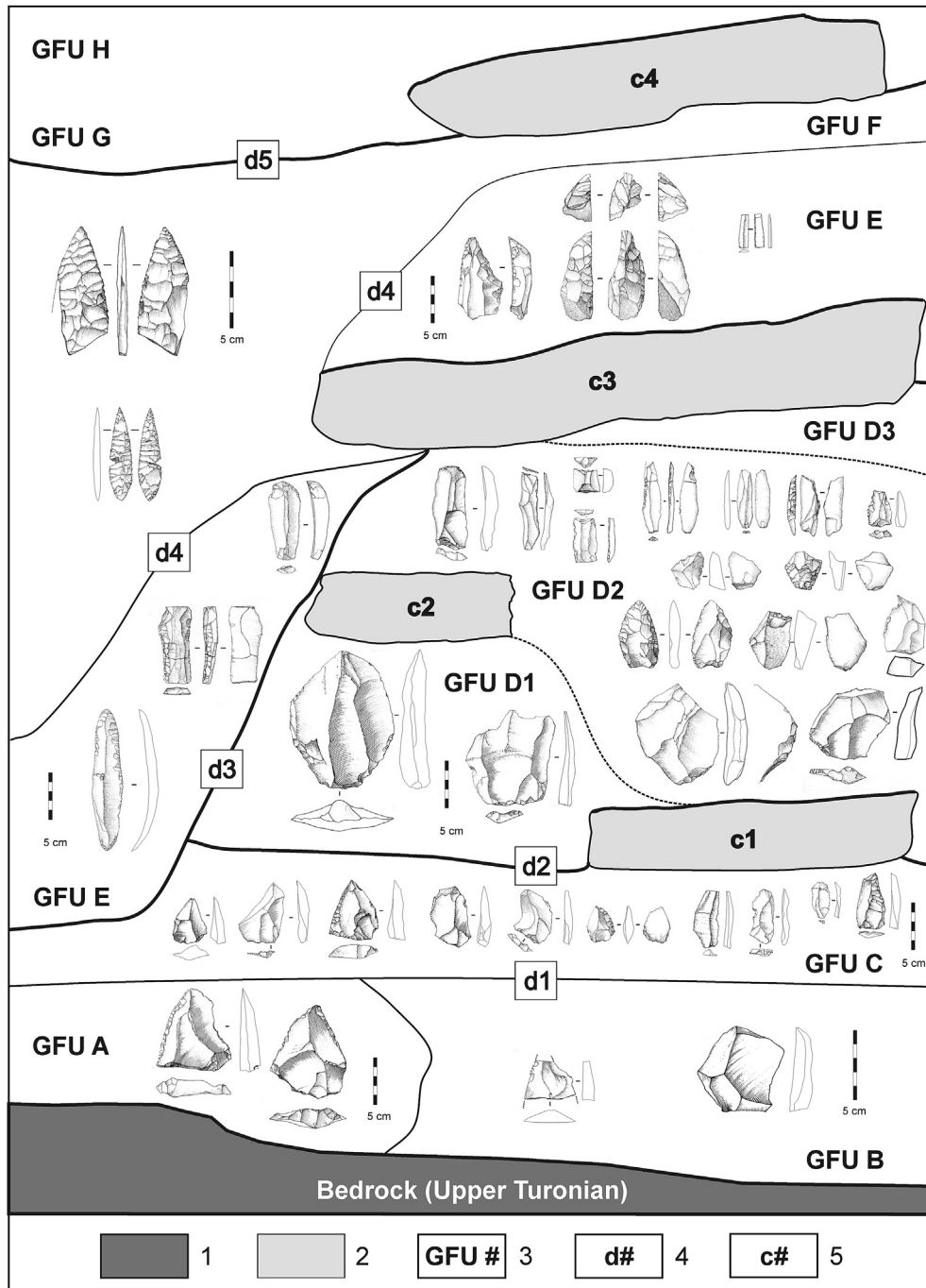


Fig. 4. Synthetic sequence of selected lithic tools recovered in the GFU A-F of the LRA-BF rockshelter. Bio-calcuridite and calcarenite (1), silicified limestone (by roof rockshelter) (2), geoarchaeological field unit (3), non-depositional unconformity and disconformity (erosional unconformity) (4), rockshelter roof and wall collapse event (5).

throughout via infilling through cryptocorrosion processes (*sensu* Nicod, 1994). The initial detrital siliciclastic allochthonous inputs (GFU B and C) are related to an alluvial environment with some lacustrine characteristics, as expressed by grain size, texture and sedimentary features (Table 1). These alternate with at least two erosive phases (d1 and d2), coupled with the first important event of cave roof collapse (c1) that results in breakdown calcareous blocks lay on the GFU C (Figs. 2 and 3). This alluvial/lacustrine environment, within a possible context of endokarst fluvial dynamic, is subsequently interrupted by several phases of roof/wall collapses near the cave/rockshelter entrance (GFU D) (*i.e.* the

Geoarchaeological Complex 2 – GC 2) (Fig. 3). This resulted in the accumulation of pebbles (GFU D2) and large boulders with volumes in the order of cubic meters together with a sandy input (Table 1). These breakdown events could suggest a phase of decompression of the rock mass, which may be the result of climatic factors, seismic events, or a rapid retreat of the cliff related to regional morphodynamic evolution. They indicate the start of a rockshelter opening, able to receive sediments from the slope. Around and under the broken-down roof blocks of the GC 2 complex, a major eroded channel (d3) was filled by an alluvial sequence that developed along the slope (GFU E and F), which was followed by a further

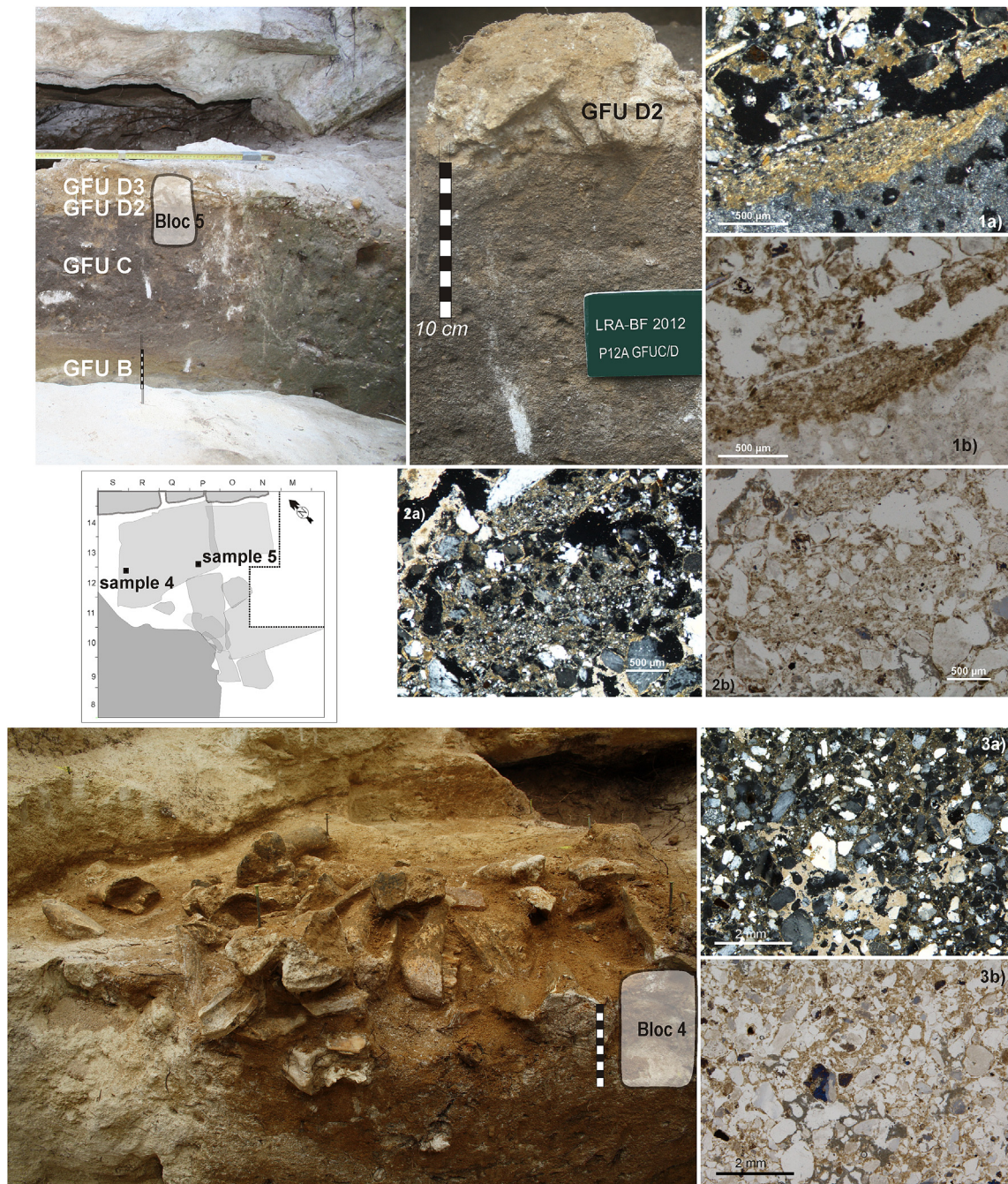


Fig. 5. Thin section positions and microphotographs related to the GFU D1 and D2. 1a and 1b: silty clay cappings on the upper part of the gravels, 2a and 2b: layered coating of unsorted silt, GFU D2: impact of melt water percolation during seasonal surface soil thaw, 3a and 3b: porphyritic related distribution of the GFU D1: water logging of the sediments during an episode of seasonal surface soil thaw.

phase of roof breakdown, and consequently the enlargement of the rockshelter entrance (GC 3 complex). Two erosive phases are evident during the filling of the eroded channel (d4) and after the last rockshelter roof collapse event (d5). The dynamics that led to the sedimentation of the upper part of the sequences (GFU G and H) (GC 4 complex) are gravity-driven, with some concurrent action of water and inputs from the slope.

4. Archaeo-stratigraphic and technological evidences

The GFU A consists of *in situ* cryptocorrosion of the Turonian bedrock, which has occurred throughout the formation of the

entire sequence, beginning prior to the deposition of the GFU B and up through the deposition of the GFU F (Fig. 3 and Table 1). The lithic assemblage recovered in the GFU A deposit, spatially restricted to the squares N9 and O10, where GFU B is absent, has yielded a small lithic assemblage (63 pieces), but some pieces are technologically diagnostic. It contains 3 large flakes obtained from large recurrent centripetal or preferential Levallois cores (Boëda, 1994; Dibble and Bar-Yosef, 1995) that have been transformed into side-scrapers through moderate retouch (Fig. 4).

The lithic assemblage recovered from GFU B is also sparse (59 pieces). Some of the flakes were produced using a recurrent centripetal Levallois scheme (Fig. 4).

Table 1
Micromorphological characteristics of the Bordes-Fitte rockshelter geoarchaeological field units F to B, and corresponding geoarchaeological complexes (GC).

GC	GFU (LRA-BF)	Micromorphological characteristics
3	F	Basic components are 90% coarse silt and a small proportion of fine sand. Abundant muscovite may correspond to transported micaflakes, or remnants of limestone weathering. The microstructure reveal an apedal material, gefuric and porphyric fabric, and vesicular porosity elongated with vertical orientation. Calcitic hypocoatings exists around the vesicular voids.
	E	basic components are 60% fine and medium sand and 40% coarse sand and gravels. The gravels are bioclastic and silicified limestone. Muscovite is represented by micaflakes of alluvial origin. The microstructure reveal an apedal material with pellicular grain structure or compact grain structure, and chitonic fabric with silt coatings around grains. Textural pedofeatures have been observed: lenticular beddings (sorted grains), silt layer up and medium sand layer down and capping on the top of free grains (silt and sorted quartz)
2	D3	Dark yellowish orange siliciclastic fine-medium sand. Without bioturbation. Wedge geometry. The predominant grains of quartz are angular and subangular with high sphericity and well sorted. The matrix is silt. The lower boundary is clear and regular. Thickness is variable, reaching up to some 40 cm.
	D2	Basics components are sandy clay loam. The microstructure reveals a chitonic and gefuric related distribution and a granular and pellicular grain structure. The micromass present a granostriated b-fabric. Chert or flint gravels and stones are presents in the ground mass. Textural pedofeatures have been observed: layered coating of unsorted silt (microphotographies 2a et 2b, fig. 5) and silty clay cappings (microphotographies 1a et 1b, fig. 5) on the up part of gravels. Secondaries carbonatations are presents (impregnation of the ground mass).
	D1	Basic components are sandy clay loam. The microstructure reveals an enaulic (chitonic or porphyric by places ; microphotographies 3a et 3b, fig. 5) related distribution and a granular and pellicular grain structure (compact grains by places). The micromass present a granostriated, a reticulate-striated and a stipple speckled b-fabric. Small fragments of bones (2 mm max.) are abundant in the ground mass, glauconites are presents, secondaries carbonatations also.
1	C	Basics components are clayey sands. Sands are heterometric. The microstructure reveal an enaulic, chitonic and by places gefuric and porphyric related distribution and a granular and pellicular grain structure. The micromass present a granostriated, a reticulate-striated and an undifferentiated b-fabrics. Some calcitic coatings exists inter-grains, associated sometimes with root tissues. Any coarse layers (medium and coarse grains, with gravels) have been noticed by places. Two types of untypical peds have been observed: one brown, rounded, with an isotropic or stipple speckled birefringence and small dark particles. The second one is a silty clay loam, leached or washed, with a lenticular or lamellar structure.
	B	Basic components are clayey sands with glauconites and carbonate gravels in places. Glauconites are abundants and have differents alteration aspects (the last one could be a clay granule). Sands are heterometric. The microstructure reveal an enaulic and chitonic related distribution and a granular and pellicular grain structure. The micromass present a granostriated and a reticulate striated b-fabric. Some calcitic coatings exists inter-grains, associated sometimes with root tissues.

The GFU C contains a lithic assemblage of 1242 pieces (Table 2), mainly composed of pieces of less than 1 cm (56%), and with a high proportion of fragments produced by fire exposure (15%). The lithic assemblage produced through a recurrent uni and bidirectional convergent or centripetal recurrent Levallois method (Boëda, 1994; Dibble and Bar-Yosef, 1995), is composed of diversified raw material from regional sources. Cores are rare, at the exception of a blank prepared to be exploited and a recurrent centripetal core. Only ten of the Levallois and non diagnostic blanks are transformed in

notches, Mousterian points or convergent side scrapers (Fig. 4, Table 2).

The GFU D lithic assemblage is more numerous (9672 pieces, Table 2). Three sub-layers have been distinguished within the GFU D during 2013 excavation (Fig. 2). Sub-layer D3, corresponding to the top of the deposit (Fig. 2), is almost sterile. Sub-layer D1 was only detected under a large rockshelter roof fragment and at its northern limits (Figs. 2 and 4, Table 1). More than 95% of the GFU D lithic assemblage was recovered in the sub-layer D2. Surface and edge damage analysis shows that its majority (67.2%) correspond to rounded, highly patinated, edge-damaged, altered by freeze-thaw non diagnostic flake fragments (Table 2, Fig. 6). A small proportion of this patinated and altered component assemblage (ca. 5%) is composed of cores and flakes, produced in the framework of Discoidal, recurrent uni and bidirectional and centripetal Levallois methods. Some of these blanks are transformed into single and double scrapers (Fig. 6). This highly weathered lithic assemblage is stratigraphically associated with siliciclastic gravels that originate from the Cenozoic units overlying the site, which is evidenced by the over-representation of 1–5 cm size fraction (Table 1). Test excavations undertook along the Roches d'Abilly cliff have revealed that such altered and technological lithic materials are systematically present in the slope deposits in secondary position. Most of this lithic component does not correspond to an occupation of the rockshelter but has probably been remobilized. A lithic component, presenting the same alteration and technological features, integrating fragments of Mousterian of Acheulean Tradition bifaces, is also present in the GFU E, associated with Aurignacian lithic remains.

The remaining part of the GFU D lithic assemblage is composed of pieces damaged by heating (3.5%), and of very distinct well preserved, non-patinated, fresh-edged lithic remains (29.3%). These reveal three different typo-technological components. The first corresponds to Levallois recurrent centripetal and possibly preferential non retouched flakes transformed into single or double side-scrapers (Fig. 7). Several varieties of the local Turonian flint have been used and no core is present. The second (181 pieces) corresponds to flakes produced by a Discoidal reduction scheme (Boëda, 1993; Thiébaud, 2007; Bodu et al., 2013). Blanks were transformed into single, double or convergent side-scrapers, notches and denticulates (Fig. 8). The third component corresponds to a blade reduction sequence (318 pieces). We interpret this as production aimed at obtaining small blades with rectilinear profiles from both prismatic and burin type cores, using soft stone hammers, that corresponds with the manufacture of blanks for Châtelperronian (CP) points, which have been recovered at the site (Figs. 9 and 11). A few of the larger prismatic blades and configuration flakes were retouched into end-scrapers or truncations. The prismatic bidirectional hard hammer removal and flake-core reduction scheme found in the GFU D, and at the bottom of layer GFU E, are very similar to those described for the CP material from the French sites of Roc-de-Combe, la Côte (Fig. 1, Pelegrin, 1995), Canaule II (Bachelierie, 2011; Bachelierie et al., 2008), Le Basté (Bachelierie, 2011), La Grande Roche (Lévêque et al., 1994; Roussel, 2011), La Grotte du Renne (Bodu, 1990; Connet, 2002; Gouedo, 1990) and in Northern Spain at Labeko Koba (Rios-Garaizara et al., 2012).

Layer GFU E, which corresponds to an alluvial channel-fill along the slope (Figs. 2–4), contains both soft hammer-struck unipolar blade and bladelet on nose, and burin type cores production, transformed into tool types (“strangled” blades, end-scrapers with invasive scalariform retouch, Dufour bladelet of the sub-type Dufour) characteristic of the Aurignacian. The two ¹⁴C ages obtained suggest the presence of two distinct Aurignacian occupations of the site (Aubry et al., 2012). Finally, a Solutrean occupation is attested by the presence of bifacial thinning flakes, bifacial

Table 2
Technological, typological and surface damage description of lithic remains from the GFU D and C. Taphonomic variables, F: fresh, R/P: rounded or patinated, H: heated; technological variables, DIS: Discoidal, LCR: Levallois centripetal recurrent, LBC: Levallois recurrent bipolar convergent, BC: Châtelperronian bipolar blade and bladelet production, ND: technologically non-diagnostic blanks.

Technology	DIS			LCR			LBC			BC			ND			Total
	F	R/P	H	F	R/P	H	F	R/P	H	F	R/P	H	F	R/P	H	
GFU D total	181	131	0	21	141	1	0	0	0	318	98	6	2323	6127	325	9672
Single side-scraper	3			10	2											
Double side-scraper	2			2												
Convergent side-scraper		1														
Notch	5	1		1						4						
Denticulate										1						
Truncation																
Truncated blade										11	1					
End-scraper										4	2					
Burin										1						
Chatelperronian point										7	1					
Retouched blade										10						
Retouched flake				1												
Total retouched	10	2	0	14	2	0				38	4	0	60	54	0	184
Non retouched	171	129	0	7	139	1				280	94	6	2263	6073	325	9488
GFU C total	2	0	0	16	30	2	3	19		0	0	0	277	709	184	1242
Single side-scraper								1								
Convergent side-scraper					1			1								
Double side-scraper																
Mousterian point						1		1								
Notch				1	1			1							1	
Denticulate															1	
Retouched	0	0	0	1	3	0	0	4	0	0	0	0	0	2	0	10
Non retouched	2	0	0	15	27	2	3	15		0	0	0	277	707	184	1232

preforms and several fragments of laurel-leaves (Fig. 4) recovered in the GFU F, which follows the deposition of a sterile stratified sand deposit, and a collapse of the western portion of the rockshelter (Fig. 3 and Table 1).

5. Identifying cultural and natural formation processes at the Bordes-Fitte rockshelter

Faunal remains recovered in the GFU A, B, C and D, primarily result from anthropogenic activities, with some evidence of carnivore and post-depositional modification. In all the faunal assemblages, bison and/or aurochs and horse dominate, followed by red deer and reindeer, wholly rhino and carnivores (Felidae, Hyenidae, Canidae and Ursidae). Attempted refitting between faunal remains fragments of the different GFU has revealed a very low refit ratio (Aubry et al., 2012).

Faunal remains, independently of GFU assignation, present different state of preservation, from chalky white appearance to highly mineralized, and most of it is fragmented. The distribution of faunal remains from GFU C, D (D1 and D2) and E, reveals a direct correlation with the limits of the different rockshelter roof fragments (Fig. 10a), that could explain their differential spatial preservation.

As we have seen before, the fresh lithic remains recovered in the GFU D1 and D2 are composed of distinct typo-technological and surface damage groups (Table 2). The distribution of lithic remains and the stratigraphical position of the rockshelter roof and wall fragments reveal that the Levallois component, exclusive in the GFU D1, is mostly preserved under a large roof fragment (Fig. 10b), corresponding to the c2 event, and that Discoidal and CP lithic artifacts, recovered in the GFU D2, are distributed around this fragment (Figs. 11 and 12). Therefore, Discoidal and CP assemblages are

younger than the collapse event c2 that occurred after the deposition of GFU D1 and before GFU D2 (Figs. 3 and 4).

Technological diversity in a lithic assemblage recovered from a single litho-stratigraphical unit could be interpreted as a high resolution record of a transitional cultural process (evolution, adaptation, acculturation), post-depositional inter-level mixing (Rigaud, 1996), or the result of a cumulative palimpsest. Based on the assumption that the pre-taphonomic assemblages were originally completely non-overlapping in their typo-technological attributes, the nature and extent of post-depositional inter-level mixing affecting deposits, and their respective archaeological components, can be evaluated using different mathematical methods (Brantingham et al., 2007; Caron et al., 2011). Nevertheless, such analyses are not able to determine if an assemblage results from an accumulative palimpsest (Bailey, 2007; Villa, 1982) nor test the reliability of a “transitional” typo-technological characterization of a lithic assemblages recovered from a specific archaeo-stratigraphic unit.

In order to assess formation processes and examine stratigraphic coherence and potential disturbance of archaeological assemblages, other methods have been developed, such as the study of the orientations of elongated lithic and bone remains within a deposit, known as fabric analysis (Bertran et al., 1997; Lenoble and Bertran, 2004; McPherron, 2005). The method is based on the assumption that patterning in object orientation tends to indicate if deposits have been subject to some sort of post-depositional alteration. Statistical analysis of orientation measurements, compared to actual re-orientation pattern observed in slopes, are used to reconstruct natural processes responsible for the formation of deposits containing archaeological remains (i.e. rock-falls, solifluction, debris flows, dry grain flows, frost-coated clast flows and run-offs). Calculation derived from the position of elongated flint and bones artifacts in both GFU E and F of LRA-BF have generated

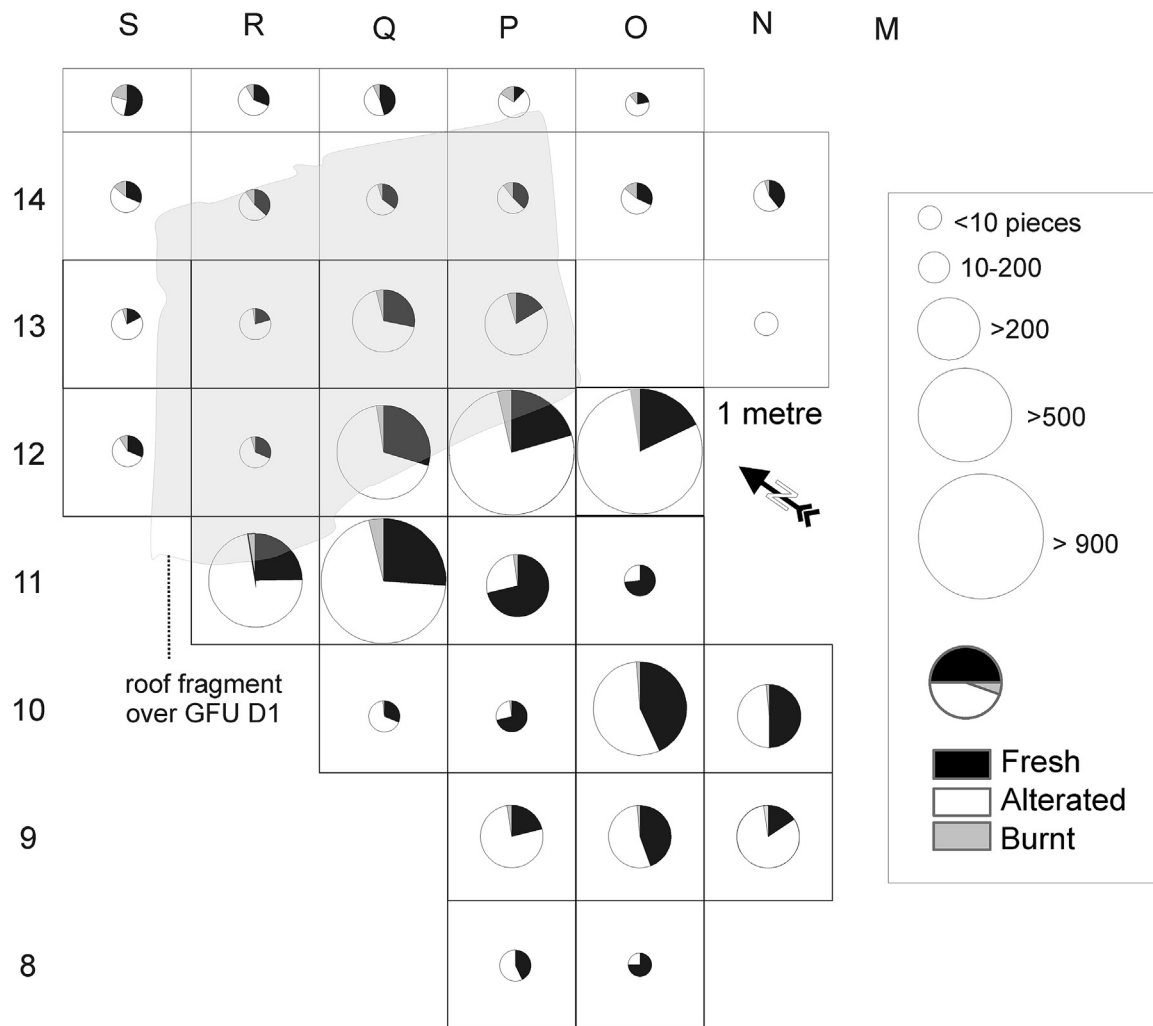
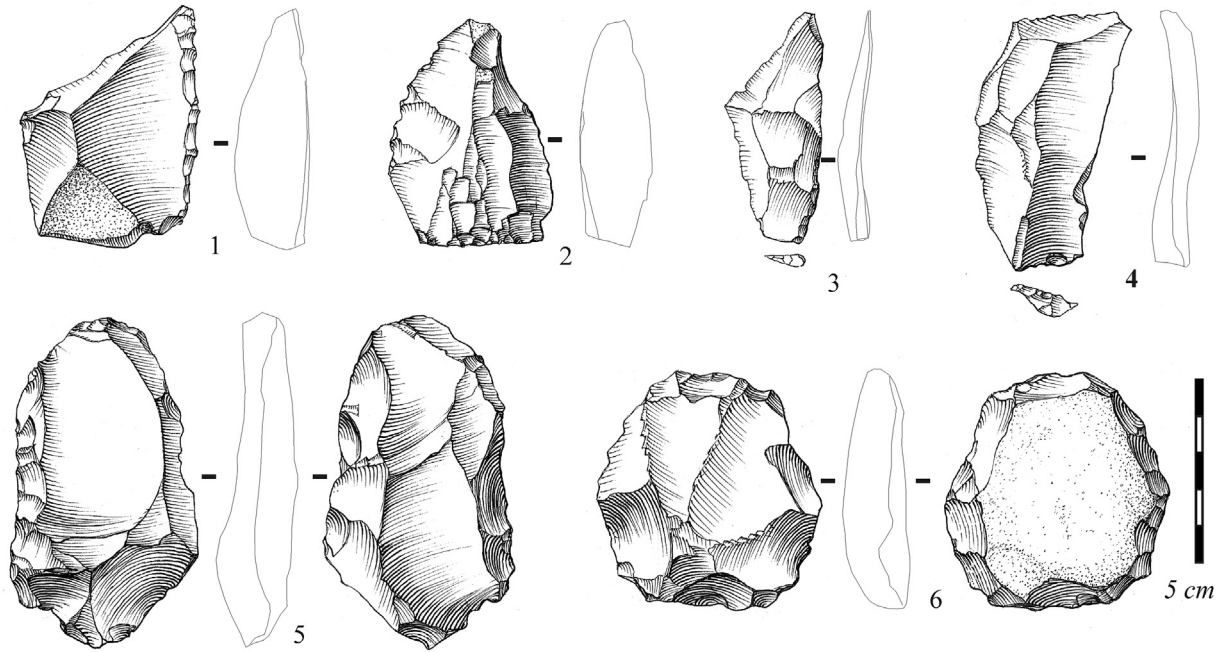


Fig. 6. Patinated, edge-damaged, and altered by freeze-thaw lithic assemblage from the LRA-BF – GFU D2, single straight (1), single convex (2, 3), double straight side-scraper (4), Levallois recurrent preferential (5) and Levallois recurrent centripetal cores (6). Relative frequency between fresh and surface damaged (patinated and rounded) remains in LRA-BF, GFU D.

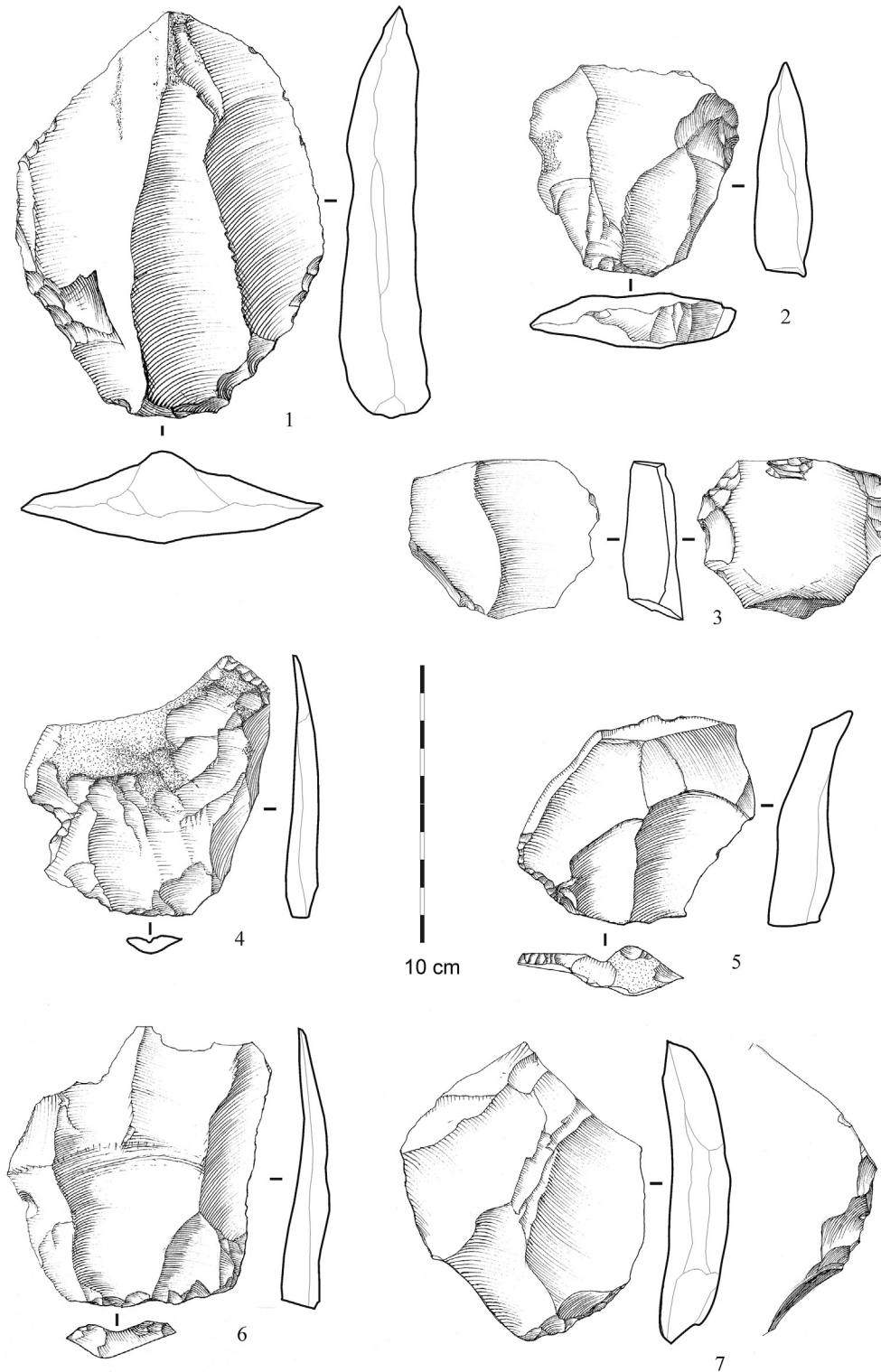


Fig. 7. Non retouched flakes obtained by a Levallois recurrent centripetal and possibly preferential method (2, 6), transformed into single convex (4, 5, 7), double convex side-scrapers (1, 3) from the GFU D1 and bottom of GFU D2.

values that correspond to a superficial run-off process, which confirms our interpretation derived just from geological observations (Table 1 and Fig. 3). That being said, fabric analysis also assumes that an archaeological sample recovered from a lithostratigraphical unit is the result of a unique sedimentary process, and do not allow precise estimation of the degree of perturbation.

The study of weathering and displacement damage combined with technological analysis have been used to argue for the coherence of lithic assemblages from layers Ejob inf and Ejob sup of the Roche à Pierrot at Saint-Césaire (Soressi, 2011). Ejob inf is attributed to a Middle Palaeolithic Levallois occupation, but the analyses are inconclusive in terms of explaining the association between the MP component and CP blade productions in Ejob sup.

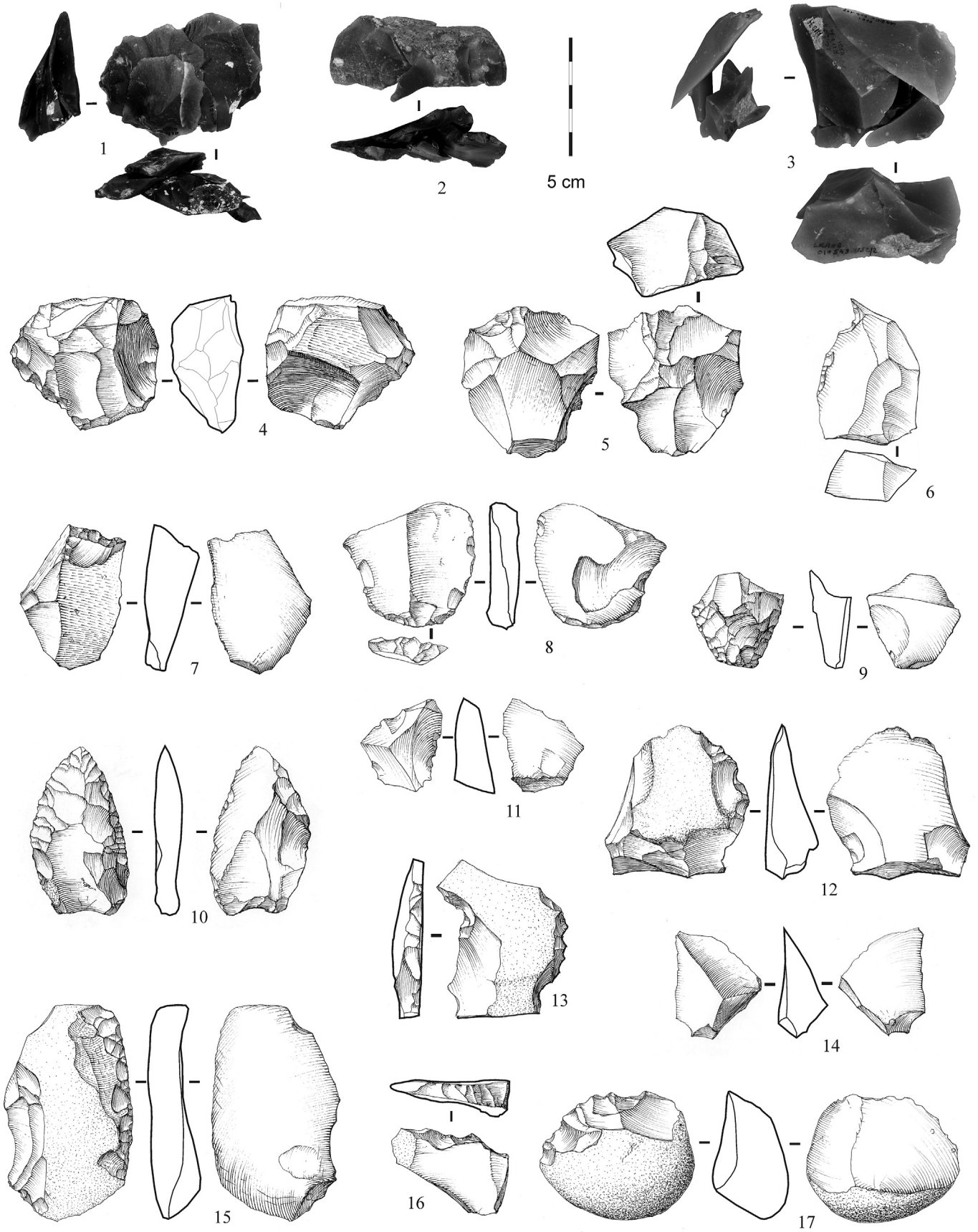


Fig. 8. Well preserved, non-patinated, fresh-edged and surface lithic assemblage from the LRA-BF, GFU D2. Discoidal flake removal refitted sequences (1–3), Discoidal reduction scheme core (4, 5), pseudo-Levallois point (14), single convex side-scrapers (15, 17), double convex, convex-concave side-scrapers (8, 9, 12, 13), convex convergent with bifacial retouch side-scraper (10), truncation (16), notch (6) and denticulate (11).

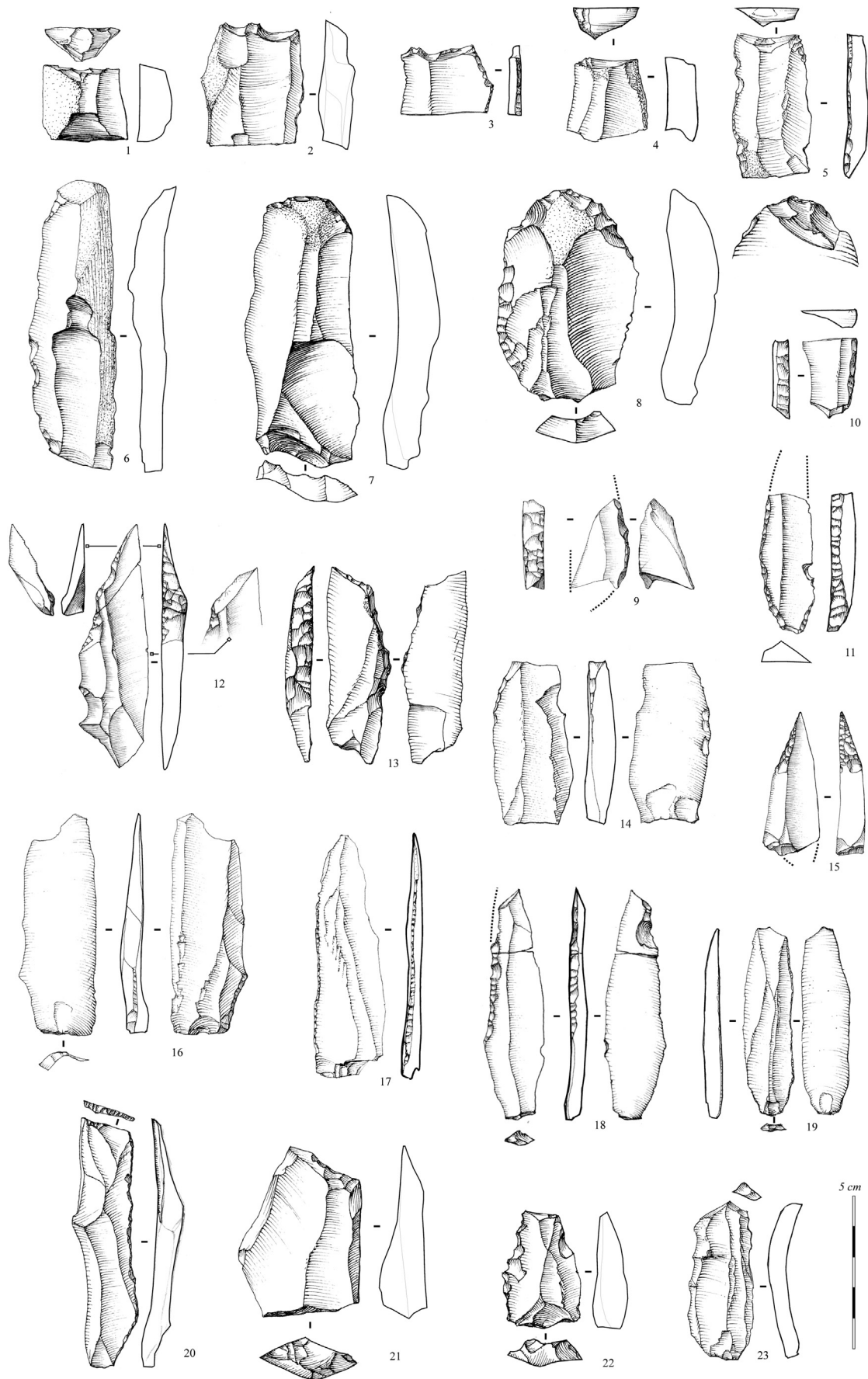


Fig. 9. Châtelperronian lithic tools, truncated blade (1–5, 21, 23), end-scrapers (6–8), Châtelperronian points (9–15), retouched blades (16–19), double side-scraper (22), and burin (20) from the LRA-BF from GFU D2 and GFU E.

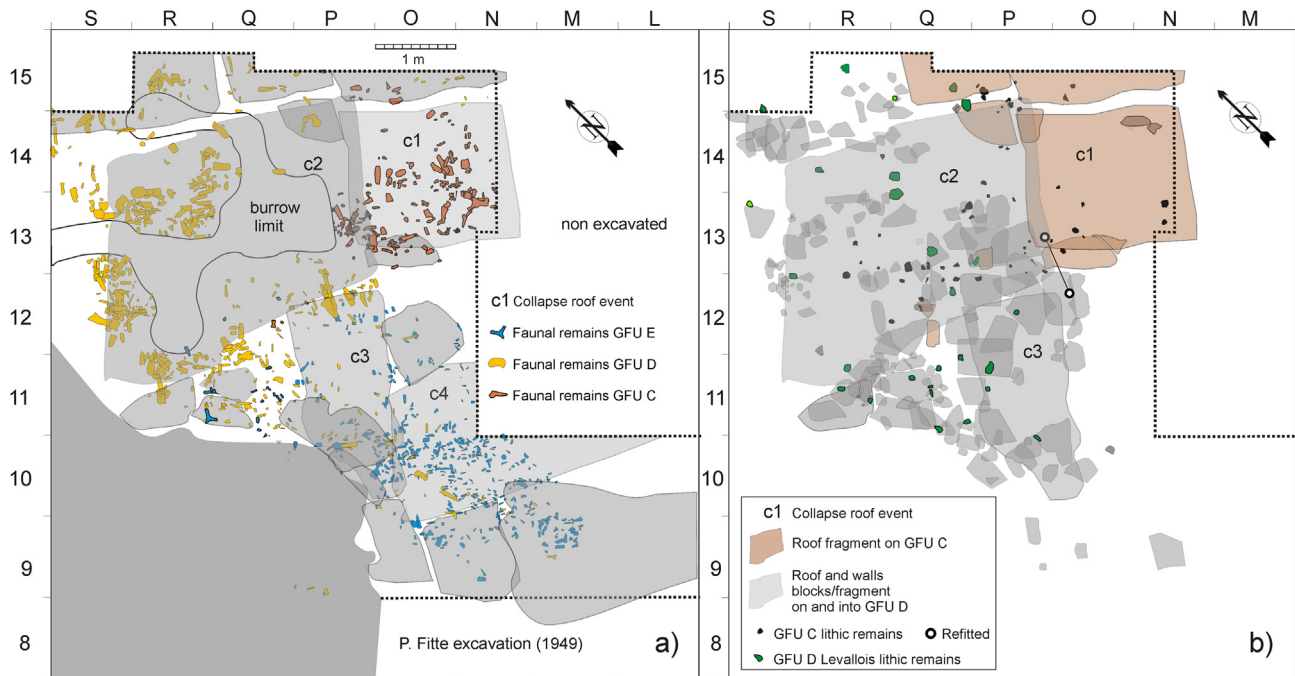


Fig. 10. Spatial distribution of faunal remains recovered in GFU C-D (a), stratigraphy of rockshelter roof fragments and spatial distribution of the lithic assemblages from the GFU C and D (b).

The study of the patinated and altered lithic assemblage of the Bordes-Fitte rockshelter's GFU D2 reveals that it is mostly located around the southern limit of a large roof fragment (Fig. 6). This component is therefore related with an increase of slope dynamics during the deposition of the GFU D2, after the roof collapse event c2 (Table 1). Such technological and surface damaged assemblage is also present in GFU E, and in all the Roches d'Abilly slope deposits (Aubry et al., 2012).

Refitting is another essential tool in assessing the integrity of archaeological assemblages (Hofman and Enloe, 1992; Discamps et al., 2012). In an attempt to alleviate the doubts concerning the integrity and homogeneity of the Ejop sup level at la Roche à Pierrot (Backer, 1993), Morin et al. (2005) carried out systematic bone refits between 8 stratigraphic faunal assemblages. In his Doctoral thesis, Bachelierie (2011) presents the results of a taphonomical and technological analysis that incorporates data from systematic refitting between fragmented flakes and blades of the Roche à Pierrot Ejop level. Bachelierie's study confirms that the Ejop inf and Ejop sup are different levels, and interprets the differential preservation, orientation, low rate of refitting (4%), and the diversity of the reduction sequences as the result of a cumulative palimpsest of Discoidal/Denticulate, Levallois/large side-scraper, and CP blade occupations. The refitting approach developed by Bodu (1990) has established the relationship between 200 pieces of the Grotte du Renne level X. Hublin et al. (2012) have recently published the spatial distribution of some of these refits in the sublayers of level X. The authors state that "the discovery of some refits between CP layers and sublayers led us to group the CP into only two chronostratigraphic units", and that "considering documented human and animal activities in the site during these two periods, considering the challenging situation of recognizing occupation floors during excavation on large surfaces of a cave deposit, and considering the limits of radiocarbon precision in this time period, further refinements in the chronostratigraphy presented here are unlikely".

Refitting has been systematically applied to LRA-BF lithic assemblages independently of their stratigraphic provenience. The lithic assemblage recovered from GFU C is essentially composed of pieces less than 1 cm (56%) (Table 2). It is made from various local flint sources and the rarity of cores corroborate what lithic industry was probably not produced on the excavated area. A unique refit was found within this level (Fig. 10b) which corresponds to a post-depositional fracture.

Refits are completely absent between the highly patinated and altered lithic assemblage of the GFU D, which are in secondary position due to the increase of slope dynamics after the rockshelter collapse event c2. For a different reason, the same applies to Levallois large flakes of the GFU D1 (Fig. 10b), which were probably not produced on the excavated area.

The Middle Palaeolithic Discoidal remains of the GFU D2 are made of several varieties of Upper Turonian flints available locally, some of which are represented by a single piece. A smaller number of Lower Turonian flint objects (20 pieces) attests to procurement from a distance of ca. 60 km. The low refit ratio for this technological assemblage (Fig. 11) could be explained by the discard of flakes or tools produced outside the rockshelter or post-depositional disturbance. These refits show a re-orientation pattern that could be the result of run-off processes in two main directions (*circa* N/S for the longest axis and E/W for the shortest) that correspond to the main dip (Figs. 11 and 13).

CP blade production is only made on Upper Turonian flint collected from Upper Turonian outcrop and fluvial terraces local sources. Refits indicate the presence of elements corresponding to all production phases for CP point production, mostly distributed by the southern limits of the large roof fragment (Fig. 11) corresponding to the collapse event c2 (Fig. 4). The pieces recovered at the northern limit (Fig. 11) correspond to the Châtelperronian points and end-scrapers blanks missing from the refit sequences. The CP production sequence yielded a high

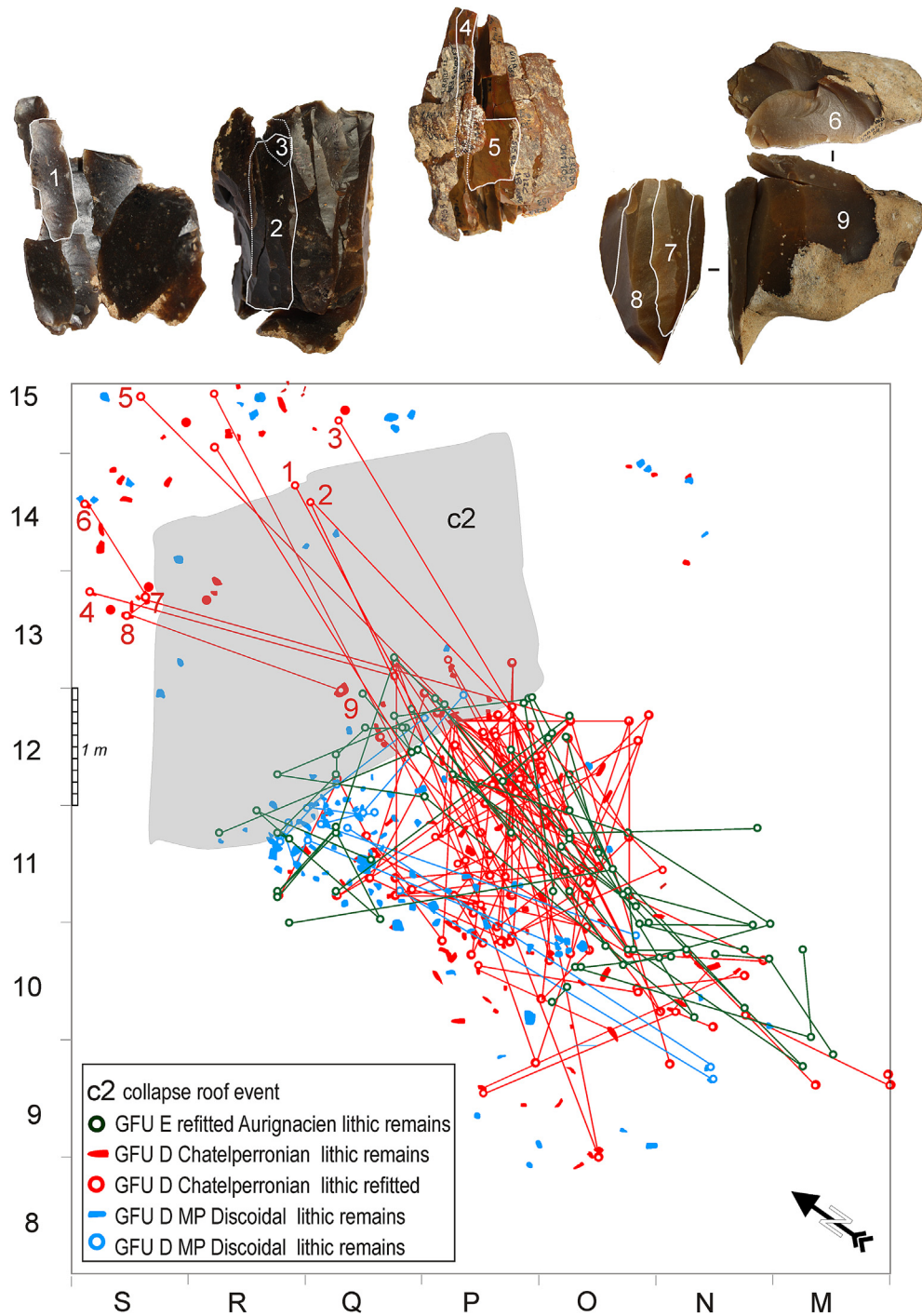


Fig. 11. Horizontal plot of Châtelperronian blade, Discoidal and Aurignacian blade and bladelet refitted removal sequences. Blades and core from 3 Châtelperronian removal sequences.

refit ratio, and most of the non refitted pieces are composed of fragment less than 1 cm. The refit rate (between 17% and 56% of the total pieces from 5 characteristic Turonian flint categories) is similar to the obtained for the refitting of experimental flake reduction sequences (Laughlin and Kelly, 2010), suggesting a low magnitude of post-depositional disturbance for the CP remains.

However, the distribution and refits of CP remains indicate that these remains were also affected by non-anthropogenic reorientations (NNE-SSW with ENE-WSW and NW-SE secondary

directions). The main direction corresponds to the orientation of refitted pieces recovered at the top of layer GFU D and integrated later at the bottom of the GFU E. These are spatially limited to the area not covered by the rockshelter roof fragments, and correspond to the pattern of orientation in the refits of Aurignacian lithic artifacts recovered in layer GFU E (Figs. 11 and 13). This result is consistent with the notion that, in areas not sealed by roof collapse, finds from the GFU D2 were removed during an erosional event (d3, Fig. 3 and Table 1), preceding the deposition of the GFU E, along an erosion channel that follows the limits of the fragmented roof

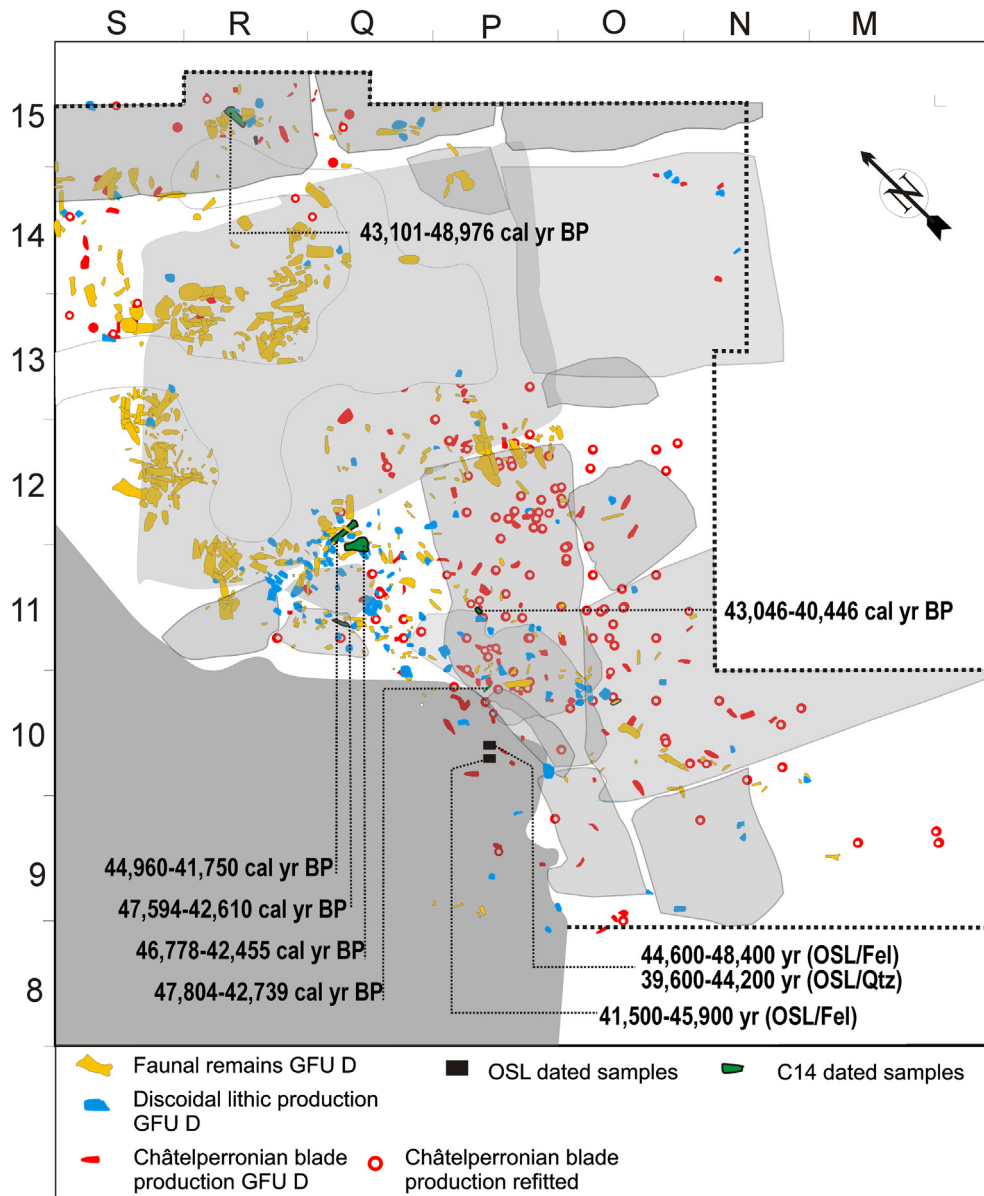


Fig. 12. Distribution of Discoidal and CP lithic remains and refitting as compared to faunal remains and bone/tooth dated by ^{14}C in the GFU D.

blocks. The secondary directions correspond to the inclination of the GFU D and to a run-off process that took place prior to the collapse event c3 (Fig. 4).

Vertical plot of the techno-typological categories defined for the lithic assemblages and faunal remains recovered in the GFU D reveals that faunal remains preserved under the largest roof collapsed fragment are stratigraphically associated with the lithic remains obtained by a Levallois *chaîne opératoire* and stratigraphically associated with the GFU D1 under the large rockshelter roof fragment or the bottom of the GFU D2 (Fig. 10). Other faunal remains, at the southern limit of the large roof collapse, are spatially associated with the Discoidal lithic remains and were affected by the same run-off process that lithic remains (Figs. 12 and 13).

The vertical projection of CP and Discoidal remains indicates that lithic artifacts corresponding to blade production overlie those that relate to Discoidal flake production (Fig. 13).

Thus, at least three human occupations characterized by distinct lithic flake and blade production could be distinguished into the GFU D, sub-layers D1 and D2 (Fig. 4). The first, defined by large recurrent centripetal or preferential Levallois blanks transformed in side scrapers, associated with faunal remains, is preserved under a large roof fragment in the GFU D1, and was affected by a run-off post depositional process around it and integrated in the GFU D2. A subsequent Discoidal occupation is concentrated at the limit of the large roof fragment in squares Q/R-11/12 and separated from the CP occupation by a wall fragment collapse event (Fig. 3).

6. Timing of Late Middle Paleolithic and Early Upper Paleolithic occupations at the Bordes-Fitte rockshelter

The (Beta-249595) age of $32,110 \pm 280$ ^{14}C yr BP (36,625–35,360 [95.4%] cal yr BP) is the only result clearly incoherent with its stratigraphic position (Fig. 3, Aubry et al., 2012). The hypothesis

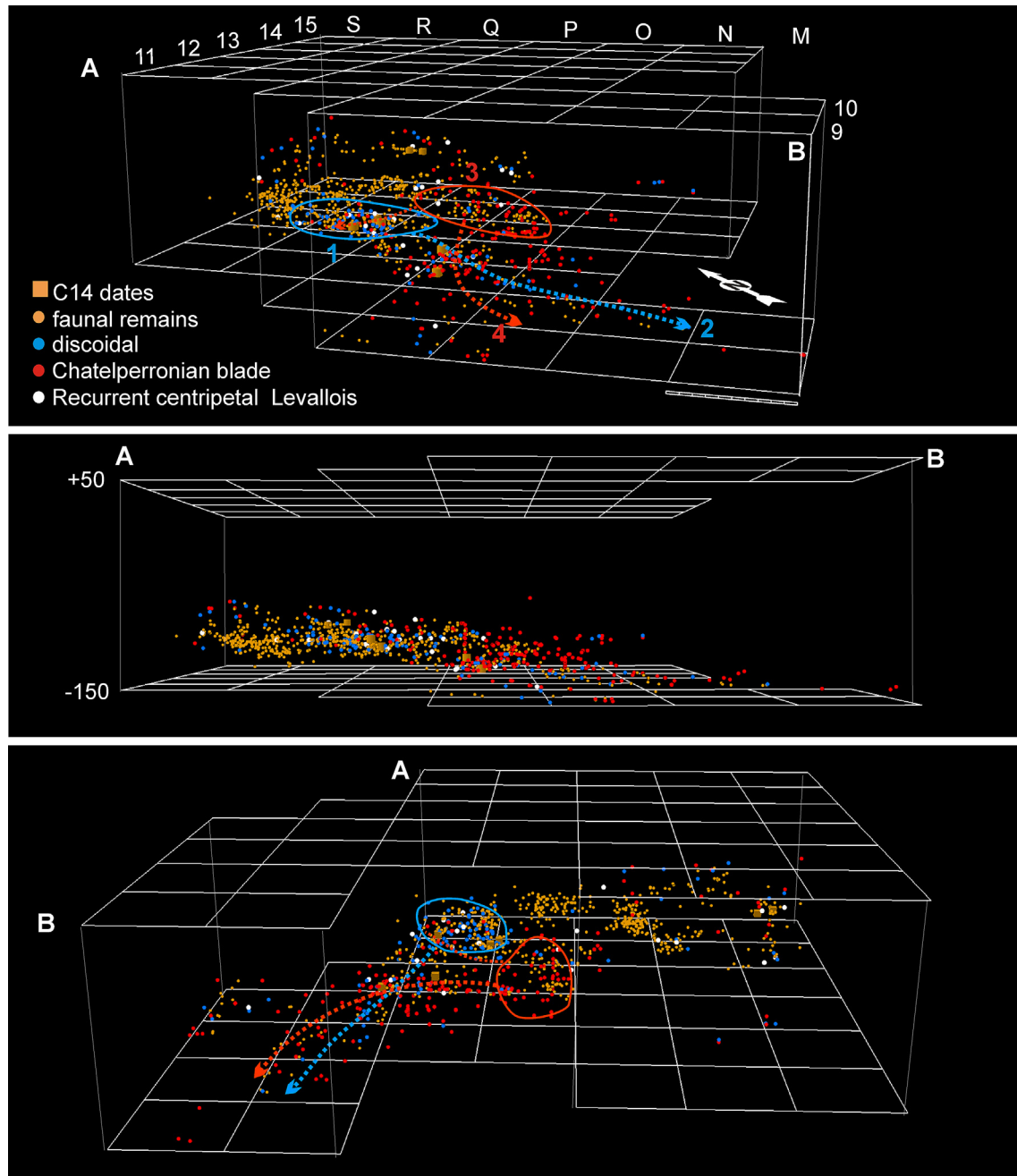


Fig. 13. Three-dimensional plotting (ArcScene 10) of the faunal remains, Levallois recurrent centripetal, Discoidal and CP lithic remains from the GFU D, and pieces from the bottom of the GFU E related by refitting (L. Luís).

more probable is of an incomplete decontamination of recent carbon during the pretreatment applying Langin protocol (Longin, 1971), confirming that most of the samples beyond 35 ka dated by this conventional pretreatment are underestimated (Higham, 2011).

The age of $41,900 \pm 1500$ ^{14}C yr BP (48,976–43,101 [95.4%] cal yr BP) obtained for a large bovid bone fragment from the GFU D1 (Table 3) gives a *terminus post quem* for the collapse event c2 that occurred between a Levallois and Discoidal occupations.

The GFU D2 ages obtained by the ultrafiltration protocol on bone and tooth are in the interval from 40.5 to 48 ka cal yr BP (Table 3). The ^{14}C calibrated ages range is consistent standard multi-grain quartz and K-rich feldspar (pIRIR₂₉₀) OSL measurements (Table 4) and additional single-grain quartz results, (Thomsen et al.,

submitted for publication). The GFU D2 samples are dated by OSL on multi-grain quartz to the 39.5–46 ka interval, and on feldspar to the 44.5–48.5 ka interval (Table 4).

The large interval for the ^{14}C ages obtained by the ultrafiltration protocol from the GFU D2 (Table 3) could be explained by their relationship with distinct human occupations of the Bordes-Fitte rockshelter (Fig. 4). The distribution of ^{14}C samples dated by the ultrafiltration protocol and lithics reveals that four bones from the 48–42 cal yr BP interval correspond to Discoidal remains, and the tooth sample dated from the 43–40.5 cal yr BP range (OxA-22342) is associated to the CP remains distribution (Figs. 12 and 13).

A *terminus ante quem* for CP occupation is provided by the ages of $34,520 \pm 850$ ^{14}C yr BP (41,095–36,925 [95.4%] cal yr BP) obtained on ultrafiltered bone collagen and $35,770 \pm 380$ ^{14}C yr BP

Table 3

Ultrafiltered gelatin radiocarbon determinations and samples that failed to produce radiocarbon results from the Bordes-Fitte rockshelter GFU D. Yield represents the weight of gelatin or ultrafiltered gelatin in milligrams, % yield is the percentage yield of extracted collagen, %C is the carbon present in the combusted gelatin, $\delta^{13}\text{C}$ are expressed in ‰ relative to Vienna Pee-Dee Belemnite, $\delta^{15}\text{N}$ values are reported relative to the AIR standard, C:N is the atomic ratio of carbon to nitrogen. Calibrated ages and boundaries are calculated by using OxCal 4.2 (Bronk Ramsey, 2013) and IntCal13 (Reimer et al., 2013).

OxA n°.	Sample reference	Stratigraphy	Material	Species	Used (mg)	Yield (mg)	%Yld	%C	d13C	d15N	C:N	^{14}C age BP	+/-	Calibrated age range (95.4% prob.)	
														From	To
26.470	LRA 11 Q12A n°63	GFU D2	Bone	Equus cabalus	690	41.12	6	44.9	-20.4	4.8	3.2	39.100	1.000	44.960	41.750
26.471	LRA 11 Q11B n°239	GFU D2	Bone	Bovid	640	27.3	4.3	44	-19.7	5.7	3.2	41.000	1.300	47.594	42.610
26.472	LRA 09 Q11A n°276	GFU D2	Bone	Equus cabalus	680	14.57	2.1	44.1	-20.2	5.4	3.2	40.600	1.200	46.778	42.455
22.316	LRA 09 P10B n°155	GFU D2	Bone	Rangifer tarandus	530	21.46	4	45.4	-19.7	5.1	3.2	41.200	1.300	47.804	42.739
22.342	LRA 09 P11A n°307	GFU D2	Tooth	Equus caballus	1050	21.96	2.1	41.7	-21.0	4.9	3.1	37.400	800	43.046	40.446
29.527	LRA 13 R15A n°35	GFU D1	Bone	Bovid	610	8.41	1.4	43.4	-20.1	5.5	3.3	41.900	1.500	48.976	43.101
	LRA 09 P10D n°110	GFU D2	Bone	Megaloceros	Failed on %N (0.4%)										
	LRA 13 R15C n°8	GFU D2	Tooth	Equus caballus	Failed due to low yield										

(41,279–39,574 [95.4%] cal yr BP), on bone collagen using the Longin method (Longin, 1971), both recovered at the bottom of the GFU E (Fig. 3, Aubry et al., 2012). Considering blade and bladelet technology of the GFU E lithic assemblage, and radiocarbon age range obtained by ultrafiltration on samples from Early Aurignacian occupations (Higham et al., 2010, 2011; Talamo et al., 2012), the date obtained on a bone using collagen extracted with alkali preparation to $35,770 \pm 380$ ^{14}C yr BP is probably underestimated (Higham et al., 2006), and could correspond to a CP occupation remains reworked in GFU E, as attested by the CP refits where GFU D is not capped by a collapsed roof fragment (Fig. 13).

Excluding the incoherent radiocarbon ages from the samples prepared by the Longin protocol and considering that multi-grain quartz 092201 and 092202 gave an estimation of the Discoidal and CP occupations, we have built a Bayesian model (Bronk Ramsey, 2009, Fig. 14) for the distribution of ages using the OxCal 4.2 (Bronk Ramsey, 2013) and the INTCAL13 curve (Reimer et al., 2013). This model, using the taphonomical and stratigraphical information, is based on the distinction of a sequence of four phases in the Bordes-Fitte rock-shelter sequence (preliminary OSL ages for the Middle Palaeolithic occupations of the GFU B and C and Solutrean radiocarbon and OSL ages of the GFU F (Aubry et al., 2012, Table 4) were not included in this model), which take into account the techno-typological data and refits. These phases are the centripetal or preferential recurrent Levallois with side-scrapers Middle Palaeolithic (GFU D1), Discoidal/Denticulate Middle Palaeolithic (GFU D2 bottom), Châtelperronian (GFU D2 top), and the Aurignacian (GFU E). The convergence value for the LRA-Bordes-Fitte rockshelter model is high (99.35) suggesting a robust model. The model obtained places the GFU D1–3 deposition in the Heinrich Stadial 5 to 4 interval (Sanchez-Goni and Harrison, 2010) and Châtelperronian/Aurignacian boundary fits perfectly with the

upper limit of the Mousterian and Châtelperronian proposed recently by Higham et al. (2014), based on AMS ages obtained by the ultrafiltration protocol from several MUP sequences of France. New radiocarbon measurements using the ultrafiltration protocol are proceeding for samples from the GFU E, in order to test the existence of two Aurignacian phases.

7. Implication for the Late Middle and Early Upper Palaeolithic timing and variability in Western Europe

Reappraising lithic assemblages from Late Middle Palaeolithic stratified sequences of Southern France, Jaubert et al. (2011) have proposed that the Mousterian of Acheulean Tradition (MTA) and CP could be separated by two LMP phases, questioning the MTA-CP direct filiation model (Bordes, 1971; Pelegrin, 1995). The first one would be characterized by Denticulate tools on Discoidal debitage, following the MTA in several sequences (Thiébaud, 2007). The second would be a “Levallois Mousterian with large scrapers”, underlying Discoidal-Denticulate assemblages, which has been proposed as a possible ending phase for the LMP in South-western France (Jaubert et al., 2011). Based on typo-technological and chrono-stratigraphy, this LMP phase has been identified at Les Rochers-de-Villeneuve, Roc de Combe, La Quina, Le Moustier, Saint Césaire, the Grotte du Bison and Galerie Schoepflin at Arcy-sur-Cure.

Levallois assemblages are overlying Discoidal levels at the top of the Middle Palaeolithic sequence of the Grotte du Bison and Galerie Schoepflin at Arcy-sur-Cure (Girard, 1982; Girard et al., 1990; Jaubert et al., 2011). Such occupations are dated on ultrafiltered bone collagen of samples from layer J at Les Rochers-de-Villeneuve to $45,200 \pm 1100$ ^{14}C yr BP (OxA-15257), from $37,600 \pm 900$ ^{14}C yr BP (OxA-2300-19), $40,600 \pm 1800$ ^{14}C yr BP (OxA-21765),

Table 4

Synthesis of OSL multi-grain ages on quartz and feldspar for the Bordes-Fitte rockshelter.

Site	GFU	Field reference	Lab code	Depth, cm	Mineral	Age, ka	Dose, Gy	(n)	Dose rate, Gy/ka	w.c. (%)
Bordes-Fitte rockshelter	F	LRA-BF 09 P12A S2	092203	170	Quartz	22.7 ± 1.3	72.2 ± 2	74	3.19 ± 0.15	4
Bordes-Fitte rockshelter	F	LRA-BF 09 P12A S2	092203	170	Feldspar	24.9 ± 1.2	100 ± 2	9	4.00 ± 0.16	4
Bordes-Fitte rockshelter	F	LRA-BF 09 N11A S3	092204	150	Quartz	26.2 ± 1.5	69 ± 2	66	2.64 ± 0.13	1
Bordes-Fitte rockshelter	F	LRA-BF 09 N11A S3	092204	150	Feldspar	28.2 ± 1.4	97 ± 3	9	3.45 ± 0.14	1
Bordes-Fitte rockshelter	D2	LRA-BF 09 P10A S4	092201	200	Quartz	43.7 ± 2.2	109 ± 2	63	2.49 ± 0.11	2
Bordes-Fitte rockshelter	D2	LRA-BF 09 P10A S1	092202	200	Quartz	41.9 ± 2.3	106 ± 3	50	2.54 ± 0.12	3
Bordes-Fitte rockshelter	D2	LRA-BF 09 P10A S1	092202	200	Feldspar	46.5 ± 1.9	156 ± 3	8	3.35 ± 0.13	3

Note: (n) denotes the number of aliquots contributing to the dose.

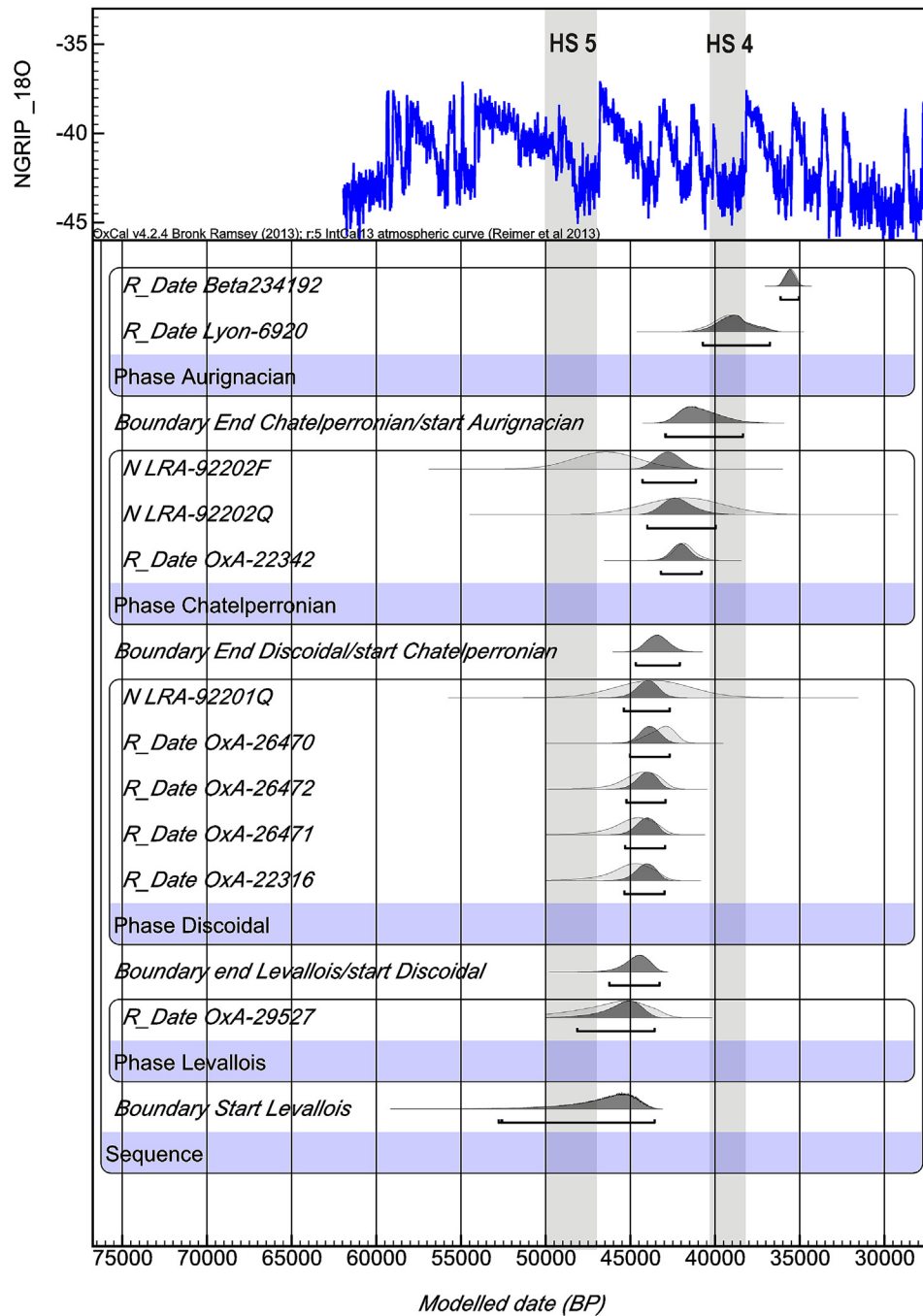


Fig. 14. Bayesian age model for the Bordes-Fitte rockshelter Late Middle Palaeolithic, Châtelperronian and Aurignacian levels produced using Bayesian software OxCal 4.2 (Bronk Ramsey, 2013) and the INTCAL13 curve (Reimer et al., 2013). The model includes radiocarbon dates and multi-grain quartz and feldspar luminescence ages from the GFU D reported in Tables 3 and 4, and radiocarbon dates published previously (Aubry et al., 2012). Shaded areas are Heinrich events 4 and 5 (Sanchez-Goni and Harrison, 2010).

$40,700 \pm 1300$ ^{14}C yr BP (OxA-2300-21) to $45,100 \pm 2300$ ^{14}C yr BP (OxA-21754) in the level J at Le Moustier, and to $41,100 \pm 1300$ ^{14}C yr BP (OxA-21805) in the level 5 at La Quina Amont (Higham et al., 2014).

At the Bordes-Fitte rockshelter, spatial analysis and refitting of lithic remains reveal that the Levallois with large side-scrapers assemblage recovered in GFU D1 is separated by a roof collapse event from the Discoidal and CP occupations of GFU D2. Stratigraphy and the age obtained for the GFU D1 do not confirm an ending position for the Levallois Mousterian with large side-scrapers of the chrono-stratigraphical model proposed by Jaubert et al. (2011). Moreover, the ages obtained by TL of 44.8 ± 3.4 ka

and 49.9 ± 3.1 ka for the Discoidal assemblage recovered at Les Bossats at Ormesson open air site (Bodu et al., 2013), the radiocarbon age obtained for the Denticulate on Discoidal debitage of Egpf level at Saint-Césaire (Higham et al., 2014) and of the level XI at the Grotte du Renne (Hublin et al., 2012), compare with the age range obtained for the Levallois with large side-scrapers occupations, do not reveal a clear chrono-stratigraphical pattern, and confirm that the LMP sequence corresponding to the Heinrich Stadial 5 to 4 interval have yet to be established.

Technological study and the lack of refits between the Discoidal and CP remains recovered from the same geoarchaeological field sub-unit D2 at LRA-BF do not provide any evidence of continuity

between the two reduction sequences. Moreover, the spatial distribution of the two technologies (Figs. 11 and 12) and 3D plotting (Fig. 13) indicate differential post-depositional displacement processes, separated by a wall fall event, and different degrees of preservation (Table 2).

The comparison with other MUP sequences dated by accelerator mass spectrometry radiocarbon on ultrafiltered bone collagen and the LRA-BF Bayesian model (Fig. 14) corroborate the age range of 45,000–41,000 cal BP for the end of the Discoidal phase and the emergence of the Châtelperronian proposed by Higham et al. (2014). However, the data also indicate that the ages associated with the Discoidal occupation of the LRA-BF are statistically equivalent to the oldest dates of the CP layers Xc-a at the Grotte du Renne (Hublin et al., 2012; Higham et al., 2010, 2014). Such a fact could be explained by large statistical error in the ages obtained for this period, or could indicate that the use Middle Palaeolithic flake-based and CP point blank technology have been contemporaneous in different regions of France and Spain (Higham et al., 2014). Another possible explanation is that the long chronological interval proposed for the CP (45–41 ka), resulting essentially from the Grotte du Renne level X ages, could result from the mixing of LMP bone remains in this CP occupations. The analysis of diagnostic lithic and bone tools distribution has shown that vertical mixing between Grotte du Renne levels is not statistically proved (Caron et al., 2011). However, the spatial pattern and inter-level refits between the sub-layers X of the Grotte du Renne present some arguments supporting complex periglacial slope deposit processes in the formation of the site (Bar-Yosef and Bordes, 2010; Bertran et al., 2010). An LMP occupation could have occurred after the erosional unconformity attested by the irregular topography of the longitudinal section of layer XI (Connet, 2002:24), and changes in the sedimentary dynamics of the layer Xbis, deposited in a humid environment (Girard et al., 1990). Reworked lithic and faunal materials could be at the origin of the Middle Palaeolithic technological component recovered in layer X and IX and correspond to the oldest dates. This human occupation, only recorded in the sequence by remains in secondary position, could be contemporaneous with the Grotte du Bison's layer G/F/E (Girard, 1982).

At la Roche à Pierrot, the other site that could indicate a transitional character for CP stone tool assemblages (Bachelierie, 2011; Soressi, 2011), species distribution and faunal refits do not show inter-level mixing (Morin et al., 2005). The direct date of $36,200 \pm 750$ ^{14}C yr BP (OxA-18099) obtained by ^{14}C using an ultrafiltration protocol on the Neanderthal tibia from Ejop sup (Hublin et al., 2012), yielding 0.77% of collagen, corresponds to the upper limit of the CP ages range proposed by Higham et al. (2014). The direct dating, contradicting the attribution of the Ejop sup occupation to an early phase of the CP (Lévêque et al., 1994), has been interpreted as an argument that the Roche à Pierrot's Ejop sup is contemporary with the most recent CP level VIII of the Grotte du Renne, and with the earliest likely modern humans remains documented in western Europe (Hublin et al., 2012). The ages obtained recently for 3 bone samples from the Ejop sup level, of $36,000 \pm 700$ ^{14}C yr BP (OxA-21699), $36,650 \pm 750$ ^{14}C yr BP (OxA-21700) and $37,200 \pm 1000$ ^{14}C yr BP (OxA-21636) are statistically equivalent to the direct date obtained on the Neanderthal tibia (Higham et al., 2014). Radiocarbon measurements and the possible existence of a burial pit attested by a lack of stones in the immediate context of the skeleton (Vandermeersch, 1993; Zilhão, 2013), are strong arguments to reject the hypothesis of an association of the Saint-Césaire Neanderthal with one of the Discoidal and Levallois lithic components attested in Ejop sup (Bachelierie, 2011), as proposed by Bar-Yosef and Bordes (2010), and to confirm the relationship of the Neanderthal skeleton with the CP lithic component of this level (Lévêque and Vandermeersch, 1980).

8. Conclusions

Sedimentary dynamics reconstruction, technological study and lithic refitting revealed that Levallois, Discoidal and CP stone tool assemblages recovered in the Bordes-Fitte rockshelter's GFU D1 and D2, result from successive occupations, separated by wall and roof collapse events, and affected by successive natural post-depositional displacement processes during the abrupt climatic oscillations of the Heinrich 5 to 4 interval (Fig. 14). Radiocarbon and luminescence ages obtained for the GFU D sub-layers, compared with other MUP sequences recently dated (Higham et al., 2014) confirm that at least a recurrent centripetal (and possibly preferential) Levallois and a Discoidal phase exist between the MTA and CP. Despite recent radiocarbon dating of Late Middle Palaeolithic occupations, the lack of taphonomical data concerning associated faunal and lithic assemblages do not permit to establish an accurate and definitive chrono-stratigraphical framework of the Middle Palaeolithic occupations post-dating the MTA.

The technological variability of some of the CP assemblages, and the 45 to 41 ka probability distribution function for the beginning and the end of this technocomplex proposed by Higham et al. (2014), could be explained by the existence of transitional industries within the same human community, the coexistence of distinct lithic technologies in the same time period, or by a bias resulting of an environmental context favorable to the formation of accumulative palimpsests.

Considering that Middle Palaeolithic flake-based technology always precedes CP blade point production when present in the same stratigraphical sequence, the technological homogeneity of all the CP lithic industries recovered in single occupation level of open-air sites, the low resolution of ^{14}C beyond 35 ka and OSL dating methods, independent age control investigation must be developed (Thomsen et al., submitted for publication). The different hypothesis should be tested applying a methodology susceptible to distinguish anthropogenic and non-anthropogenic processes in the formation of the archaeo-stratigraphic records in order to establish the exact relationship between samples dated and human or carnivore's occupation remains.

Untangling human and others contributions in site formation processes is still a crucial step regarding the elaboration of an accurate timeline and to establish the nature and cause of behavioral changes and relationship with the replacement processes of Neanderthals by modern humans in Western Europe.

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